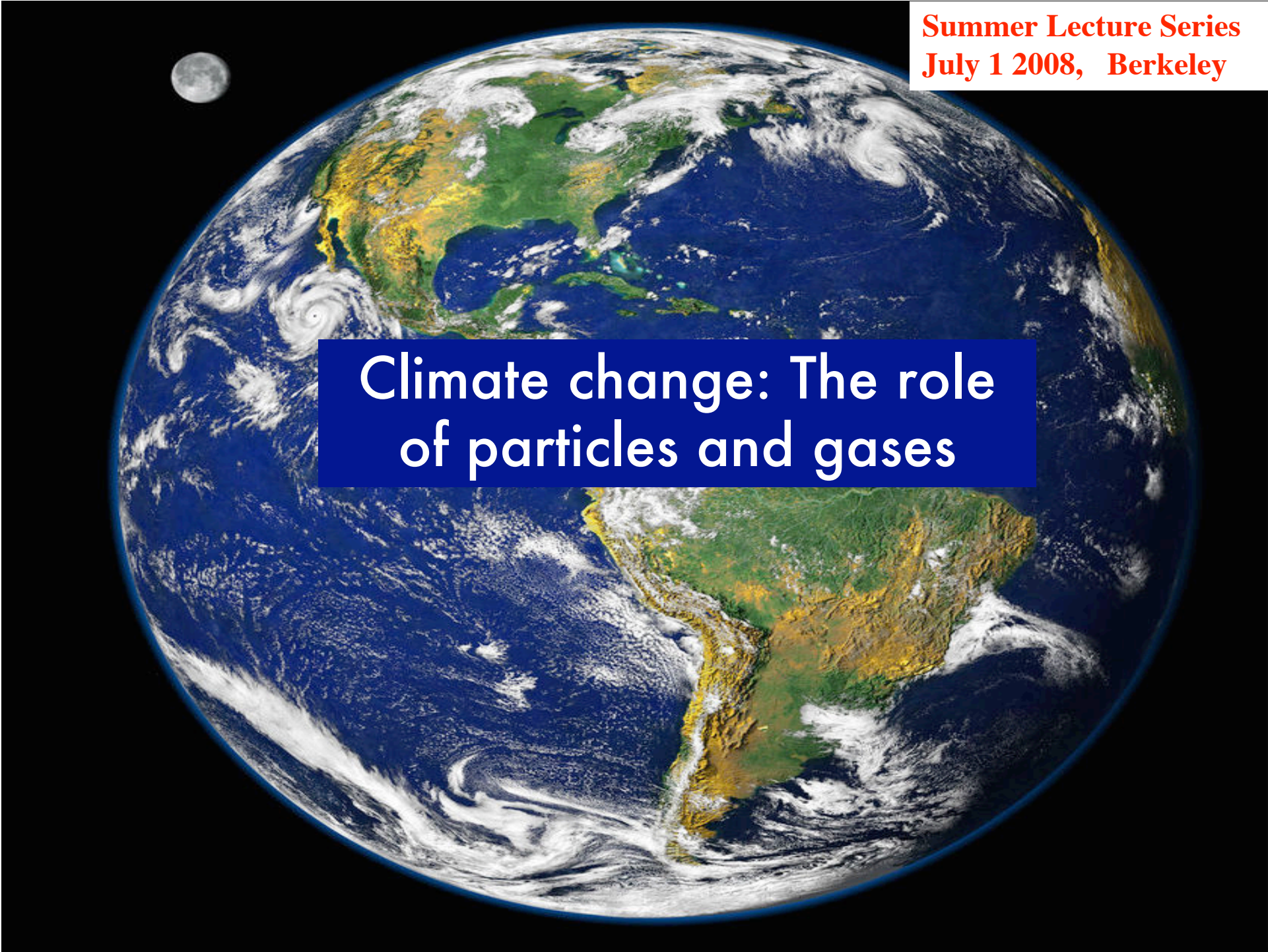


Summer Lecture Series
July 1 2008, Berkeley



**Climate change: The role
of particles and gases**

NORTH POLE ON THIN ICE

Exclusive: No ice at the North Pole

Polar scientists reveal dramatic new evidence of climate change

By Steve Connor, Science Editor

Scientists warn Arctic sea ice is melting at its fastest rate since records began:

It seems unthinkable, but *for the first time in human history, ice is on course to disappear entirely from the North Pole this year.*

(Independent, UK, June 27, 2008)

Industry/economy and the environment appear to be headed on a collision course!

OUTLINE

Climate record - ~1 M yrs to present-day

Drivers of climate change

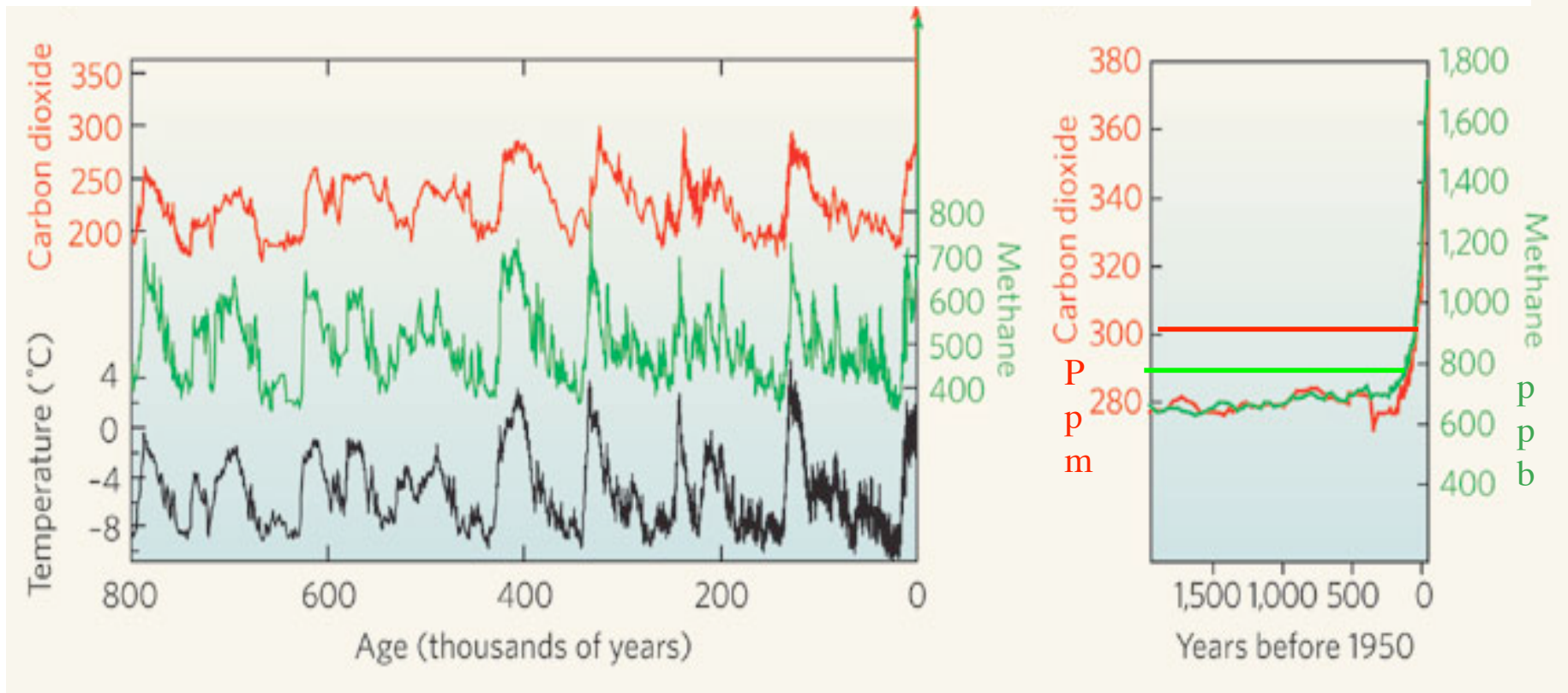
Understanding climate change

Global and Regional Signals - Industrial age

Future Climate

Earth's climate history

EPICA Dome C ice core records with a temperature reconstruction (relative to the average of the past millennium).

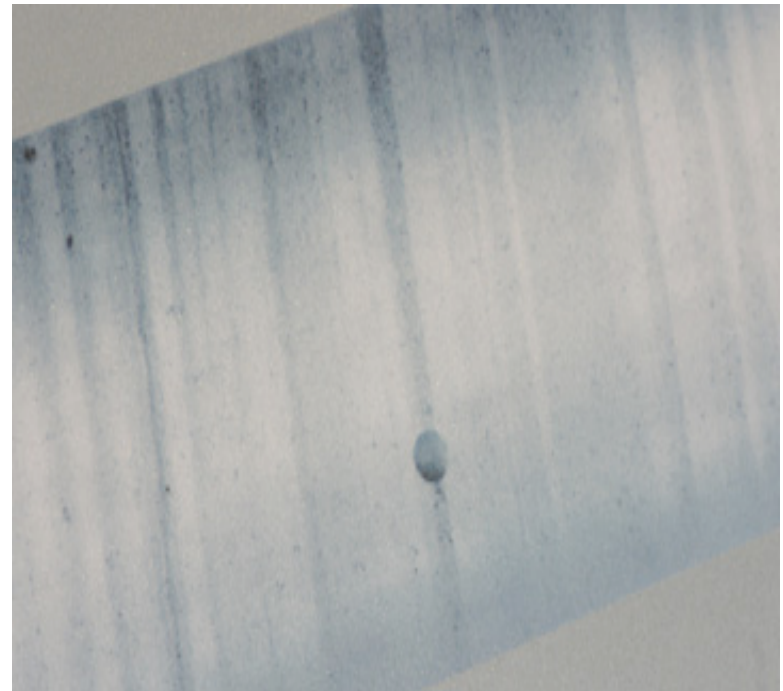


Ice cores drilled in the Antarctic ice sheet: Snowfall piles up year-by-year. Isotopic composition of the snow records the temperature and bubbles of air trapped in the ice preserve a record of atmospheric composition.

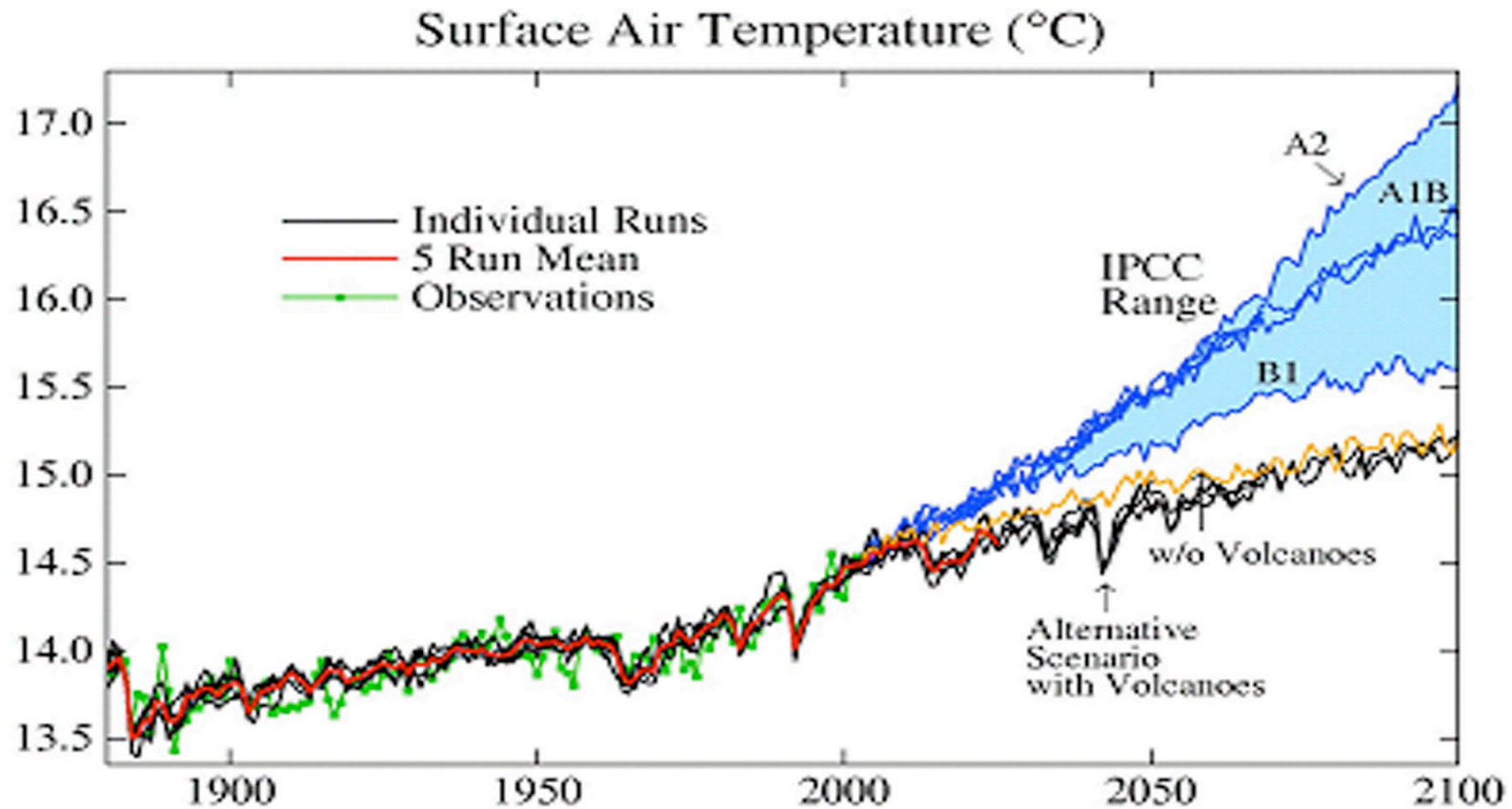
Brook (*Nature* **453**, 291-292, 2008)

Observations before instrumentation

Sediments, tree rings, ice cores, corals



21st Century Global Warming



(Hansen et al. 2007, ACP)

OUTLINE

Climate change - ~1 M yrs to present-day

Drivers of climate change

Understanding climate change

Global and Regional Signals - Industrial age

Future Climate

Natural and Anthropogenic drivers

Greenhouse gases

Carbon cycle

Nitrogen cycle

Ozone and methane chemistry

Aerosols

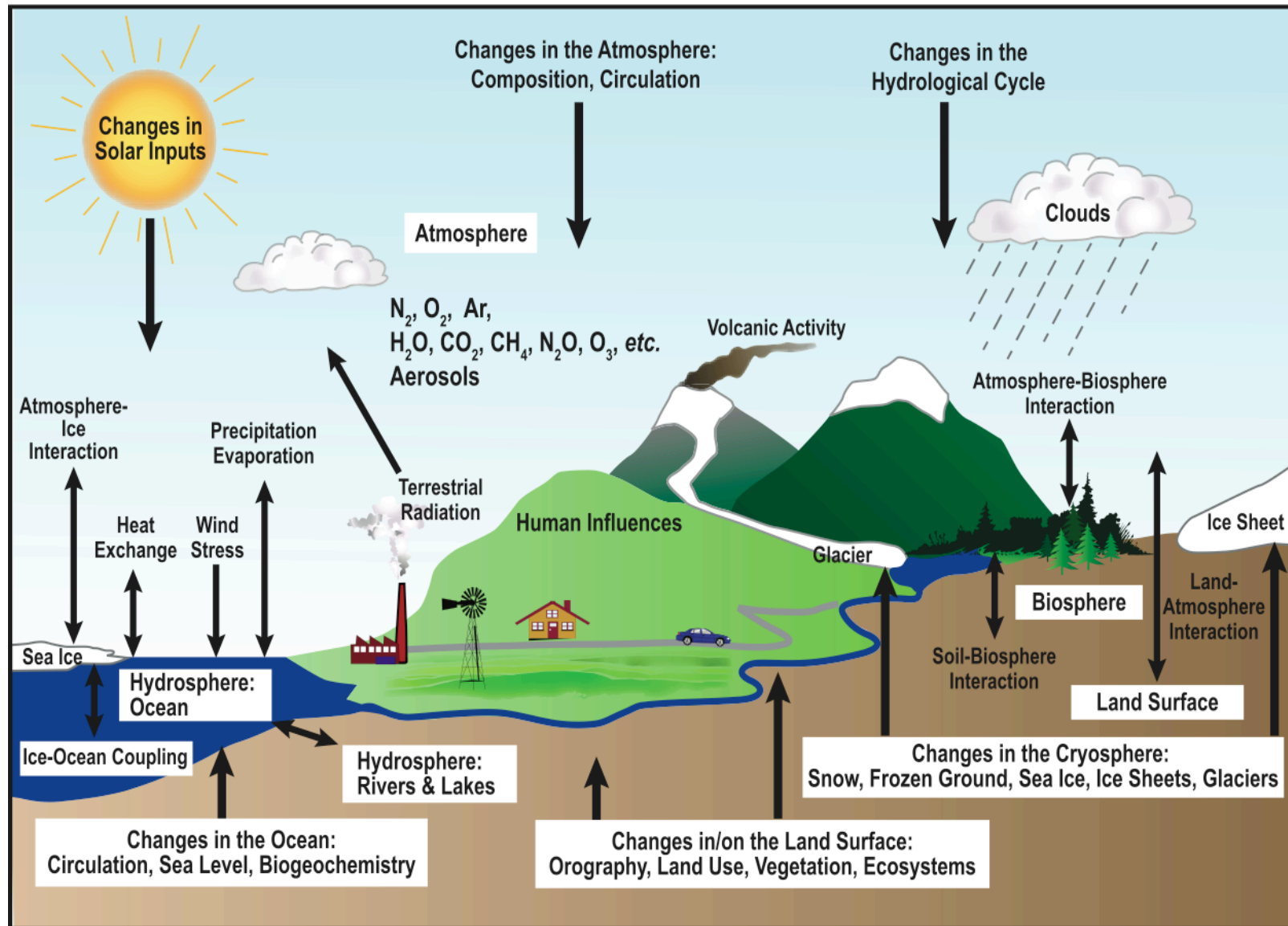
Sulfates, carbonaceous, sea-salt,
dust, nitrates

Water cycle

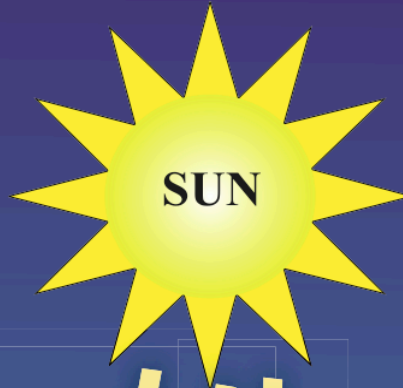
Land-use



Climate system



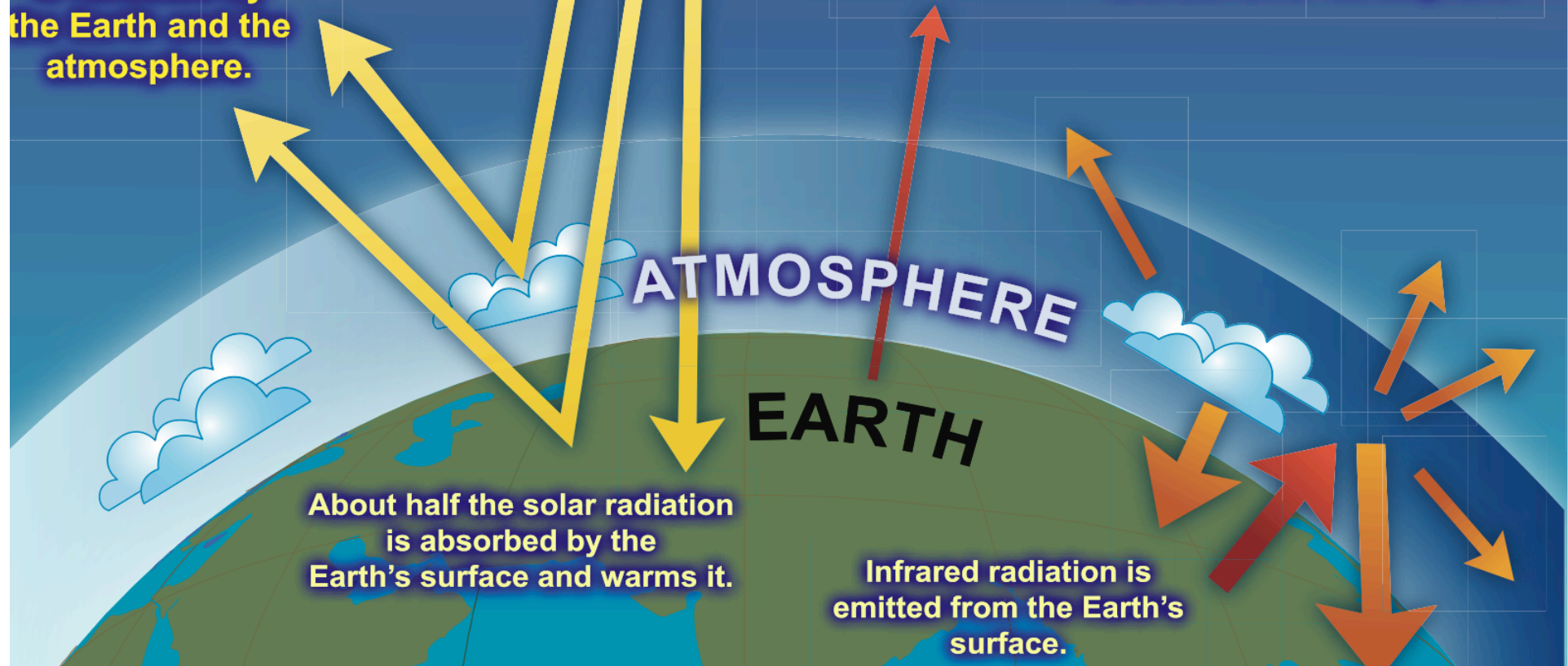
Solar radiation powers the climate system.



The Greenhouse Effect

Some of the infrared radiation passes through the atmosphere but most is absorbed and re-emitted in all directions by greenhouse gas molecules and clouds. The effect of this is to warm the Earth's surface and the lower atmosphere.

Some solar radiation is reflected by the Earth and the atmosphere.



About half the solar radiation is absorbed by the Earth's surface and warms it.

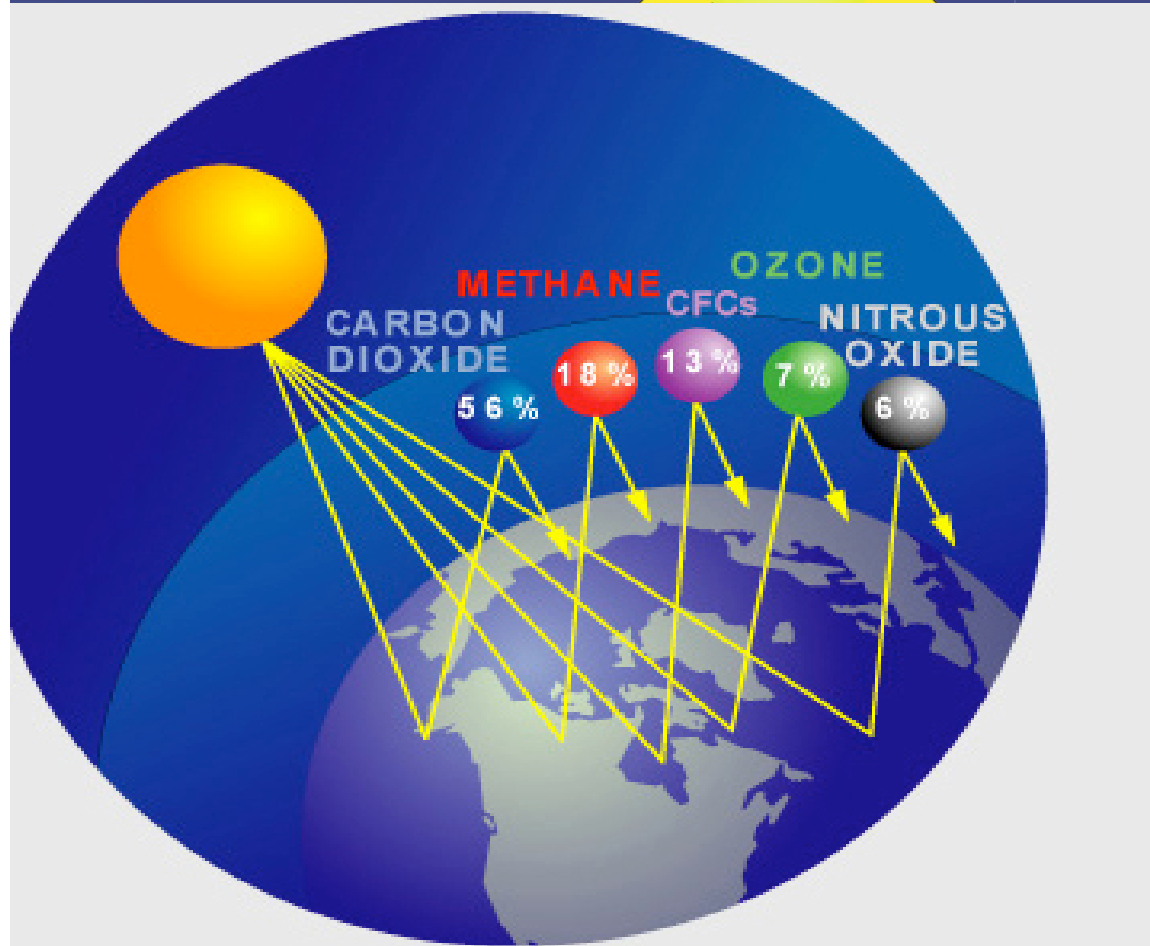
Infrared radiation is emitted from the Earth's surface.

Solar radiation powers the climate system.



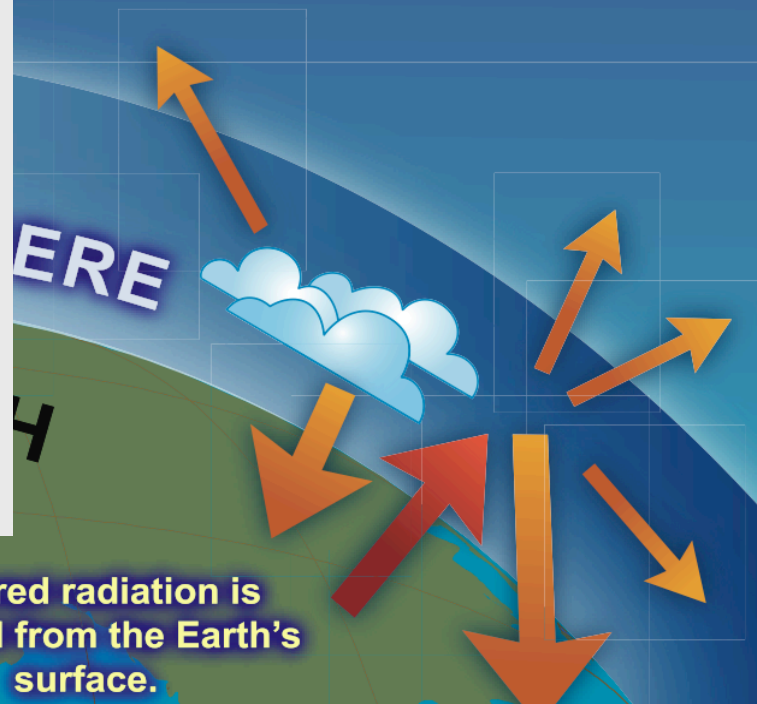
The Greenhouse Effect

Some of the infrared radiation passes through the atmosphere but most is absorbed and re-emitted in all directions by greenhouse gas molecules and clouds. The effect of this is to warm the Earth's surface and the lower atmosphere.



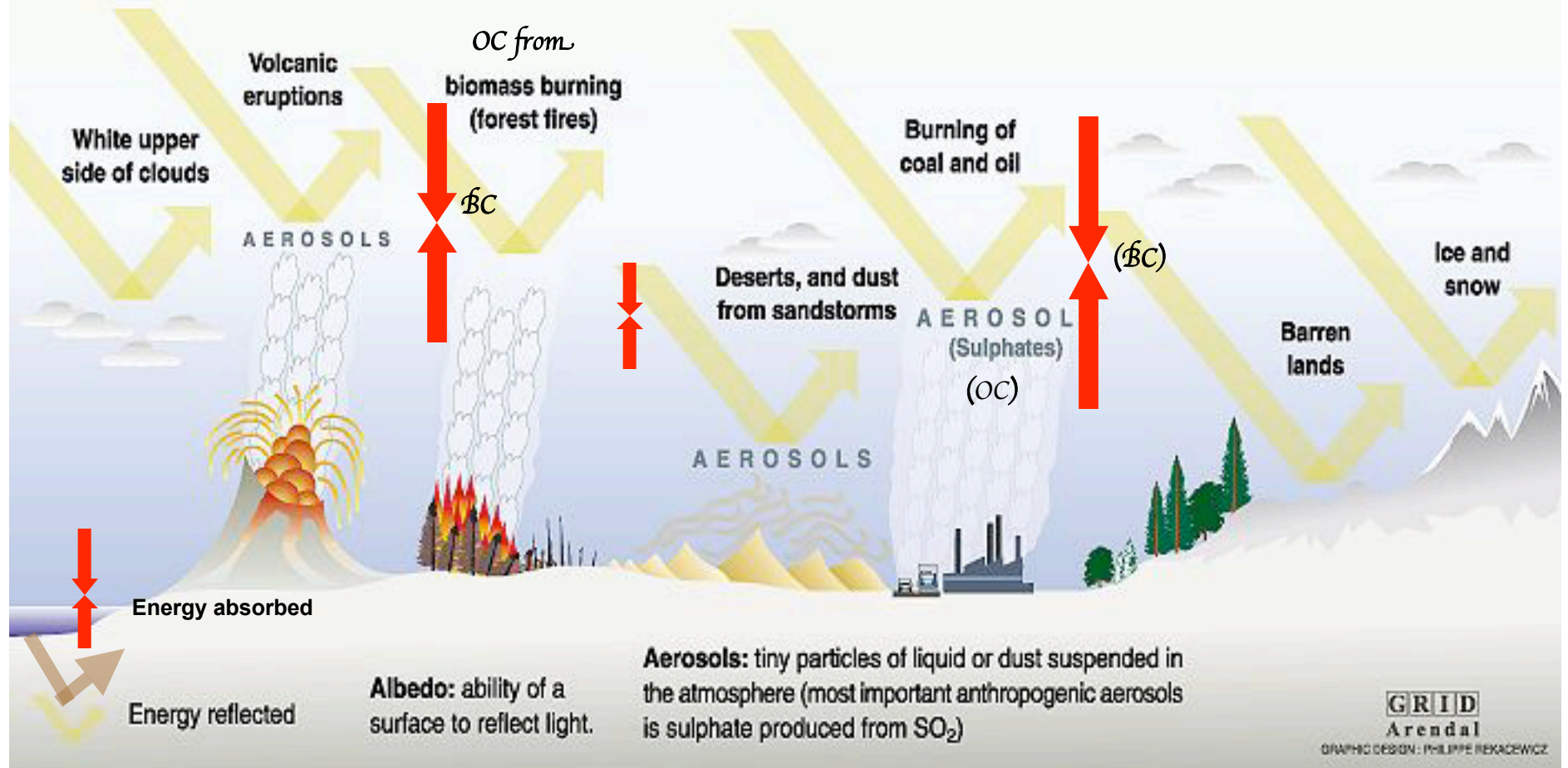
is absorbed by the Earth's surface and warms it.

Infrared radiation is emitted from the Earth's surface.



Emissions and Radiative Exchange

Aerosols - Sulfates, organic and black carbon (OC and BC), dust, sea-salt, nitrates



Sources: Radiative forcing of climate change, the 1994 report of the scientific assessment working group of IPCC, summary for policymakers, WMO, UNEP; L.D. Danny Harvey, Climate and global environmental change, Prentice Hall, Pearson Education, Harlow, United Kingdom, 2000.

Source: IPCC 2001

OUTLINE

Climate change - ~1 M yrs to present-day

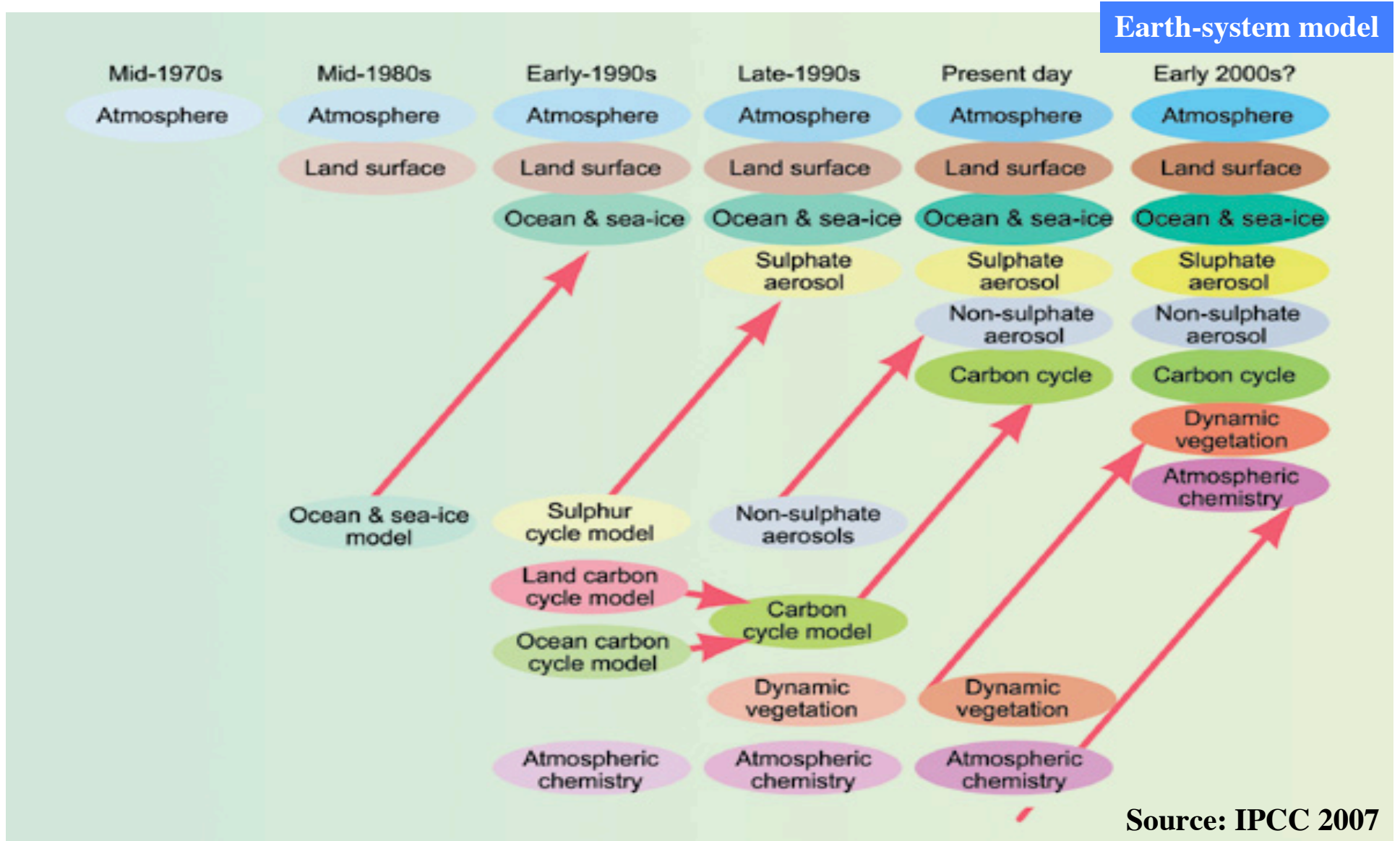
Drivers of climate change

Understanding climate change

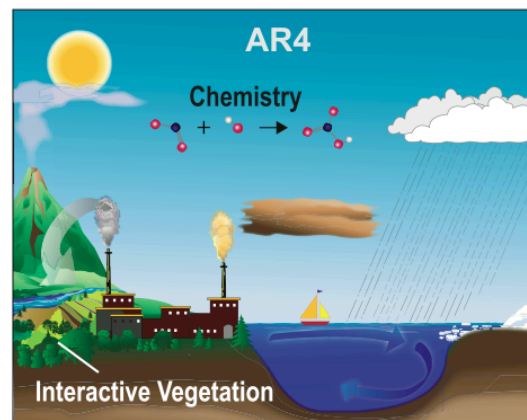
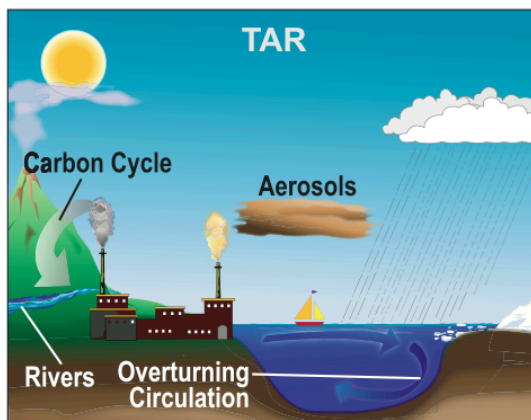
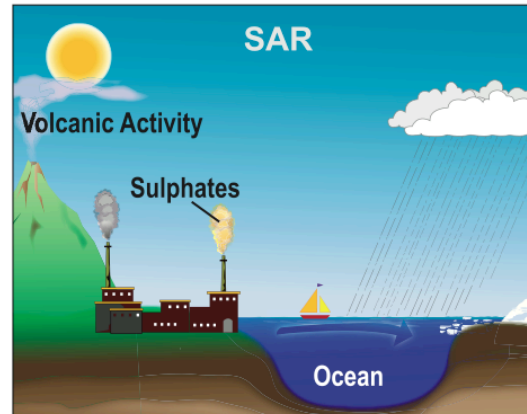
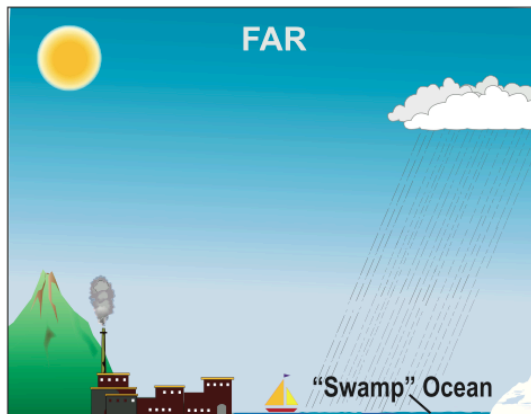
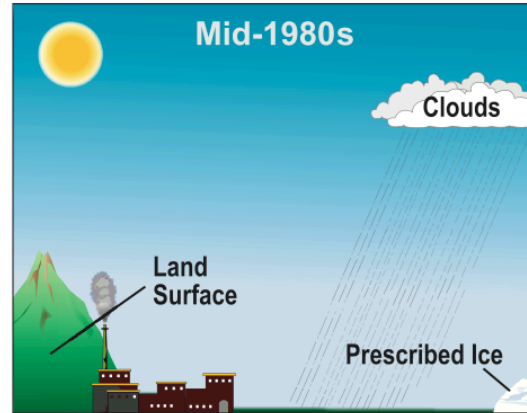
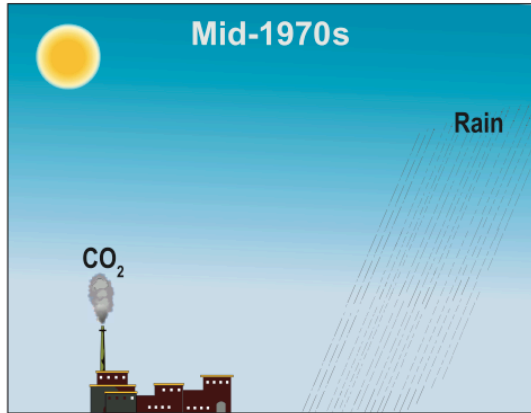
Global and Regional Signals - Industrial age

Future Climate

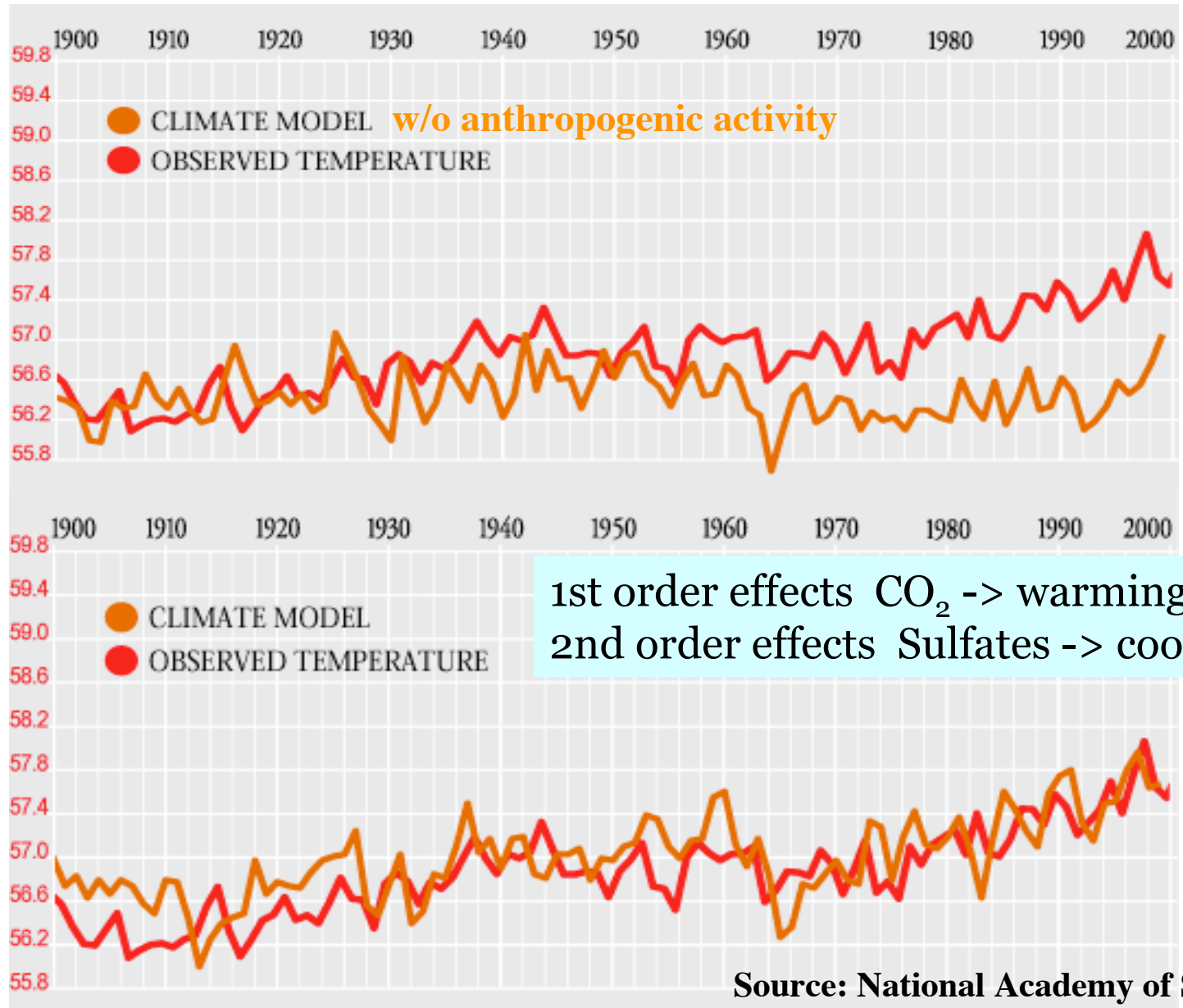
History of model development



Model landscape approaching reality



Temperature over the last 100 years

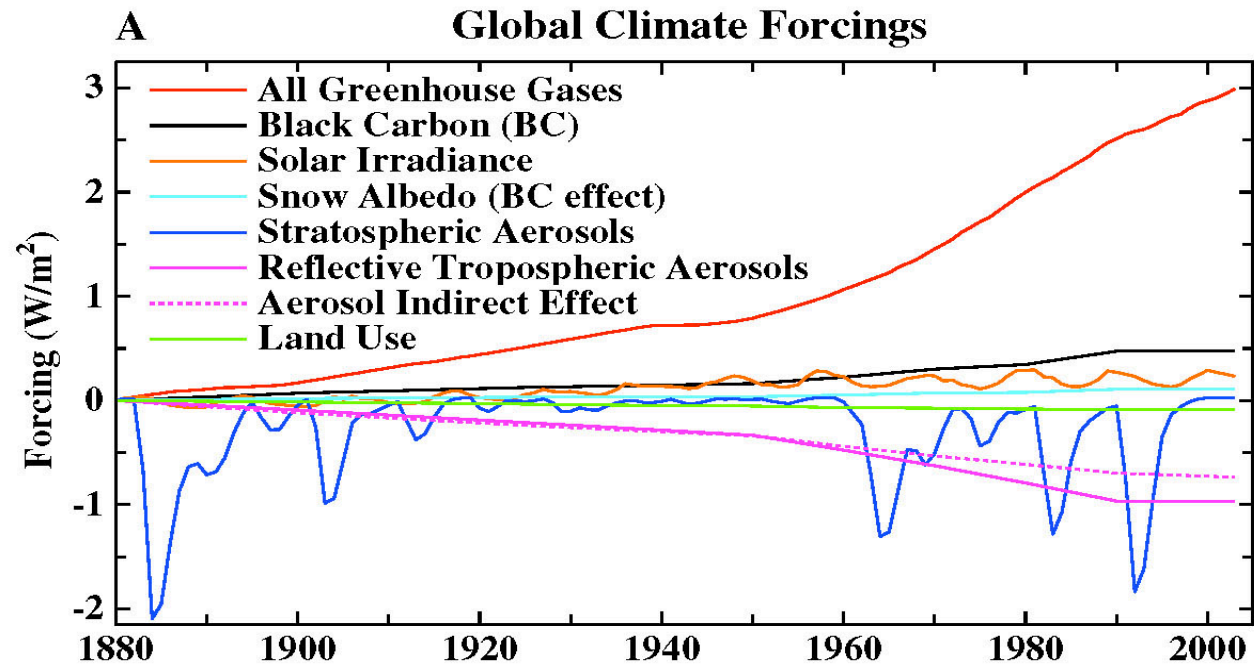


Climate model - Goddard Institute for Space Studies

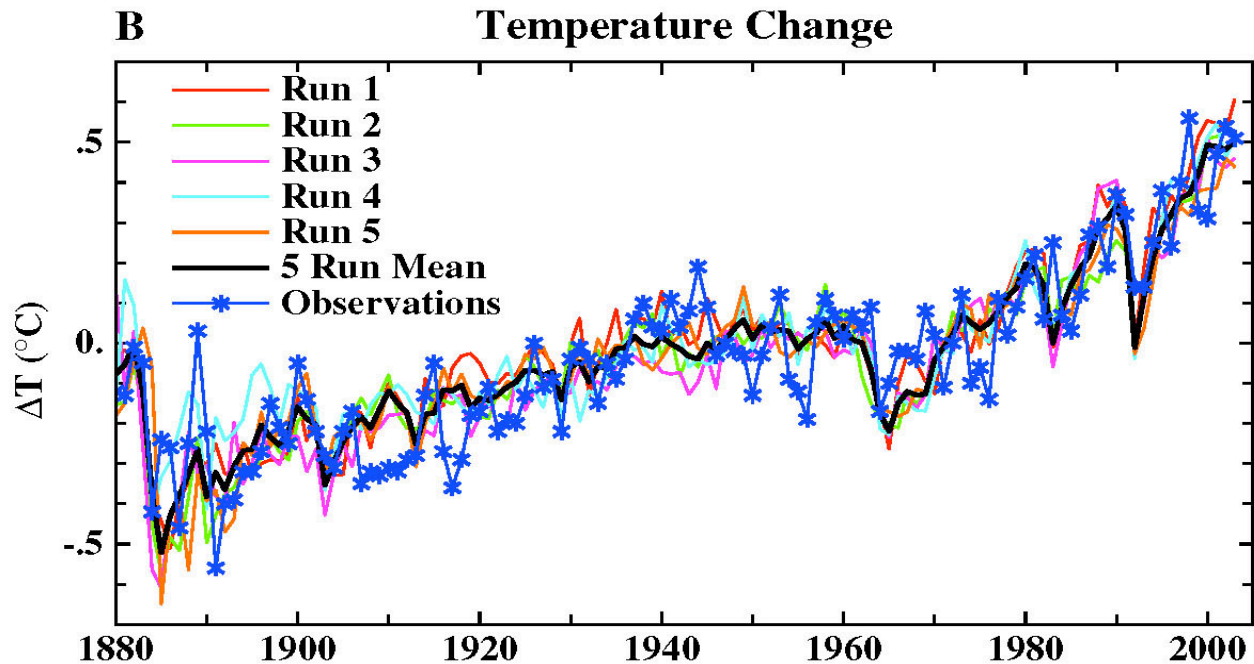
- **20-60** vertical layers, **4°x5°** horizontal resolution (Hansen et al. 1983, 2005).
- Interactive ozone, methane, aerosol chemistry (Koch et al. 2006, Shindell et al. 2007)
- Interactive aerosol-cloud-climate effects (Menon et al. 2002)
- Coupled to slab ocean or dynamic ocean/sea-ice model
- Carbon cycle and interactive vegetation being implemented

“All models are wrong, some models are useful”
(George Box, 1979)

(A) Forcings used to drive climate simulations.

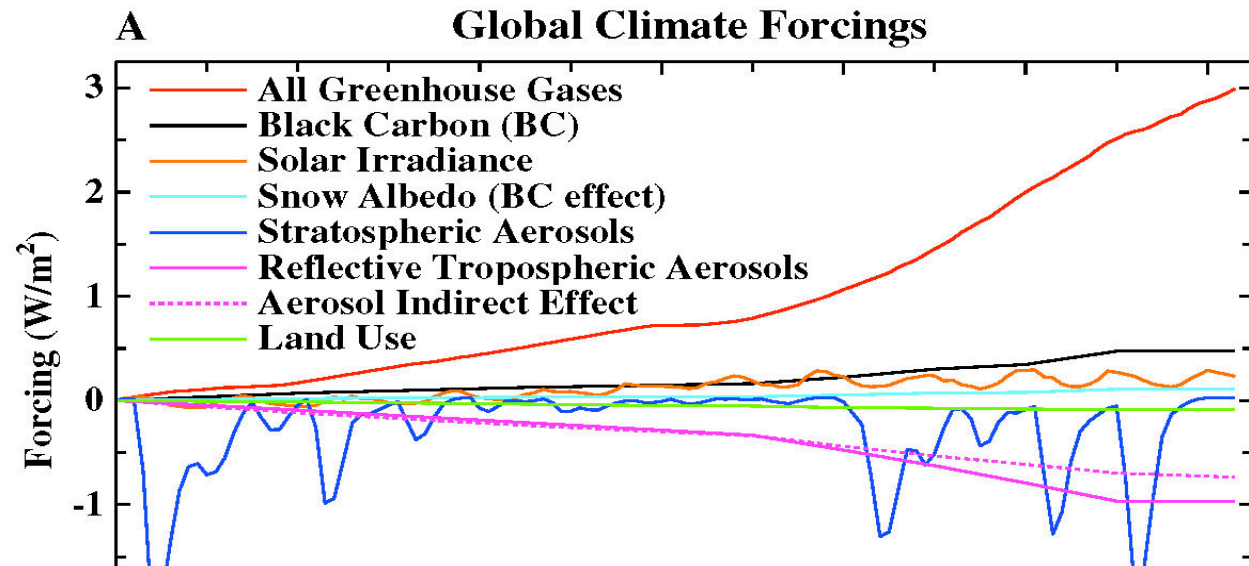


(B) Simulated and observed surface temperature change.



Source: Earth's energy imbalance: Confirmation and implications. *Science* 308, 1431, 2005.

(A) Forcings used to drive climate simulations.

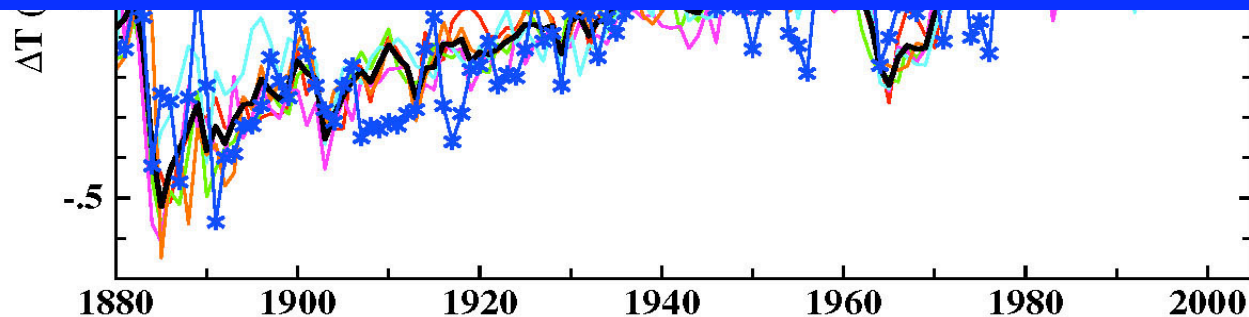


Radiative forcing: Change in the balance between energy coming in and energy going out in the climate system.

Units are W/m^2 and is usually measured at the tropopause or top of the atmosphere.

Positive implies warming, negative cooling.

change.



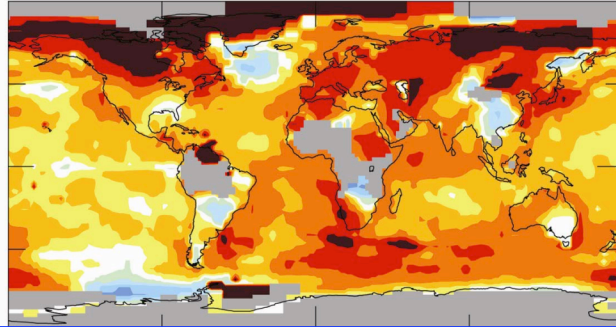
Source: Earth's energy imbalance: Confirmation and implications. *Science* 308, 1431, 2005.

1880-2003 Surface Temperature Change (°C)

Temperature change observed & simulated for different forcing mechanisms.

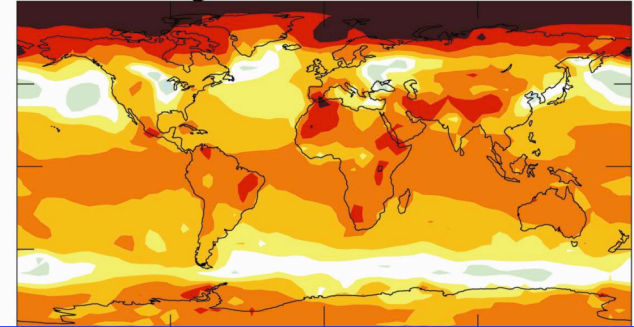
Observations

.60



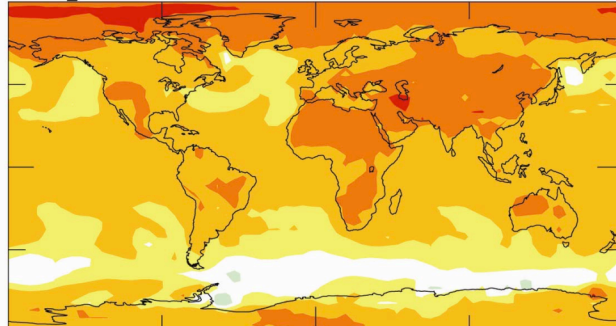
All Forcings

.55



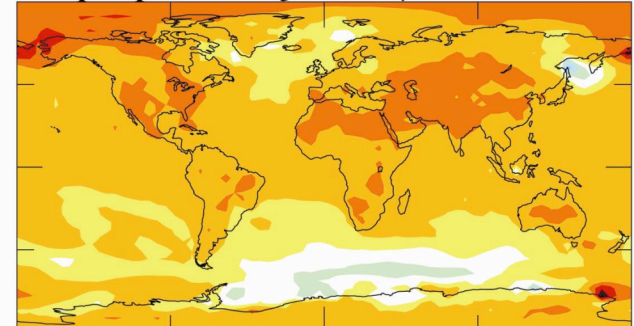
CO₂

.43



Tropospheric O₃ + CH₄

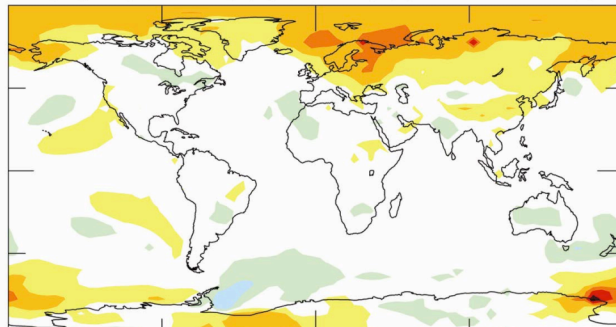
.40



Aerosol forcing (negative) is thought to be slightly excessive in in the 'all forcing' simulation.

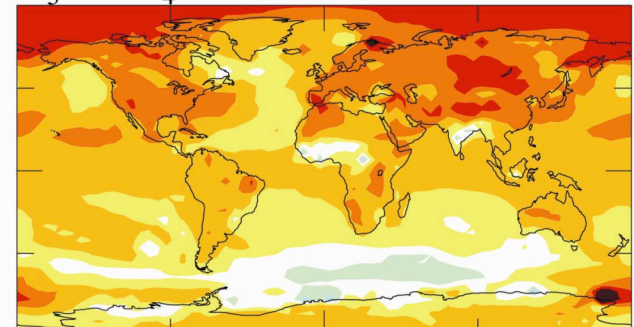
BC Snow Albedo Effect

.04



O₃ + CH₄ + BC + OC

.42

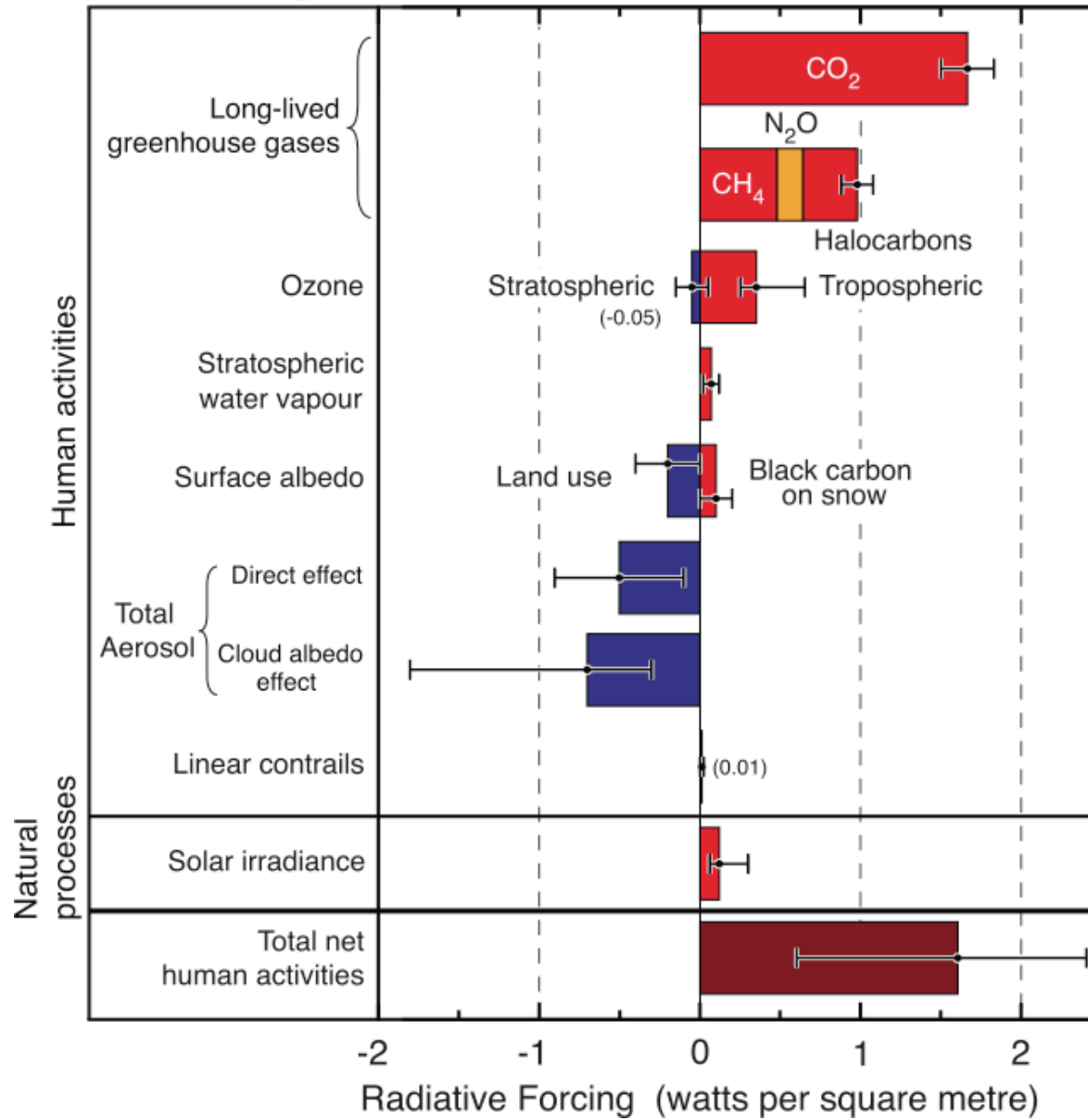


Source: Hansen et al.,
J. Geophys. Res.,
2005.



Radiative forcing of climate between 1750 and 2005

Radiative Forcing Terms



Source: IPCC 2007

OUTLINE

Climate change - ~1 M yrs to present-day

Drivers of climate change

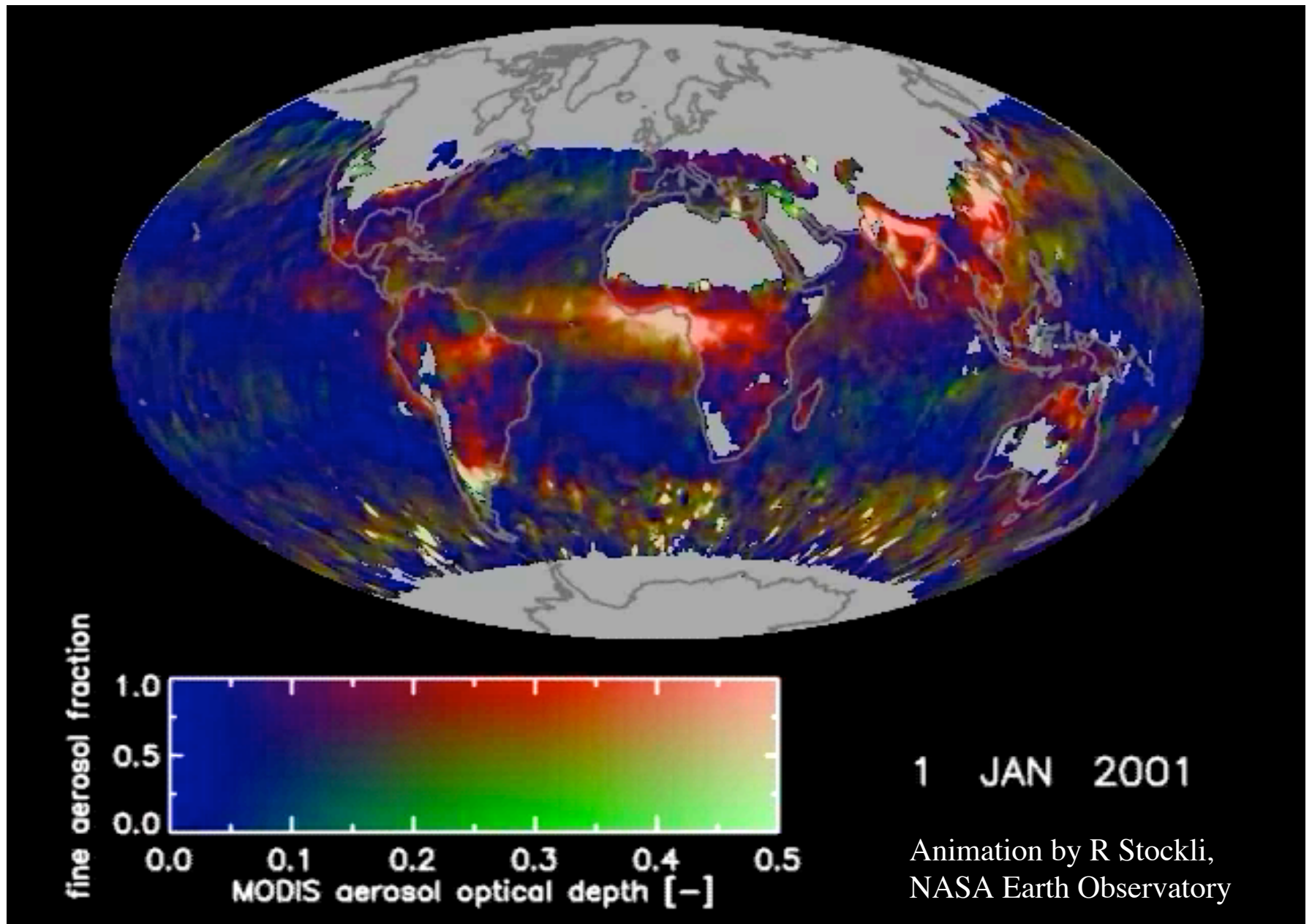
Understanding climate change

Global and Regional Signals - Industrial age

(3 examples: **GHG+Aerosols**, **Aerosols**, **GHG effects**)

Future Climate

Aerosols: Satellite optical extinction data



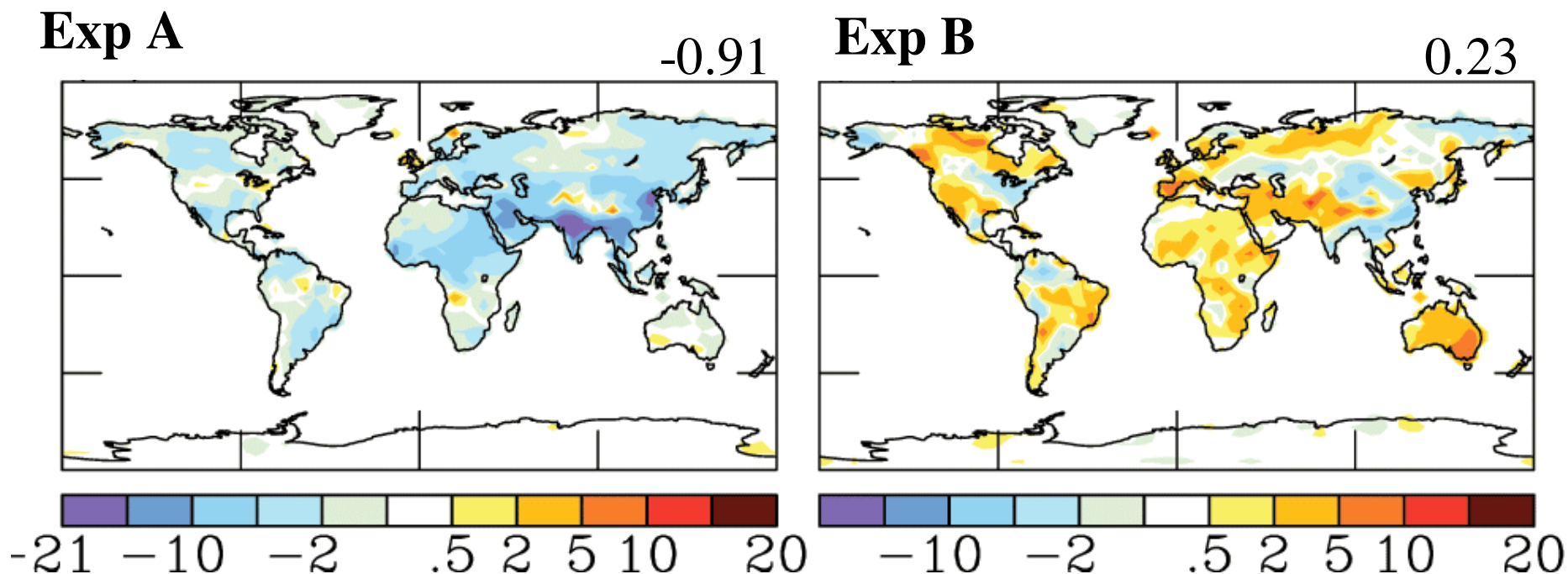
Anthropogenic Aerosol Effects: 1960-2002

Observed: “Global dimming: 1960-1990; Reversal after 1990”

We calculate linear trend in absorbed solar radiation for 1960 to 2002:

Exp A: All forcings (ozone, land-use, snow/ice albedo change, solar, GHG, water vapor, aerosols)

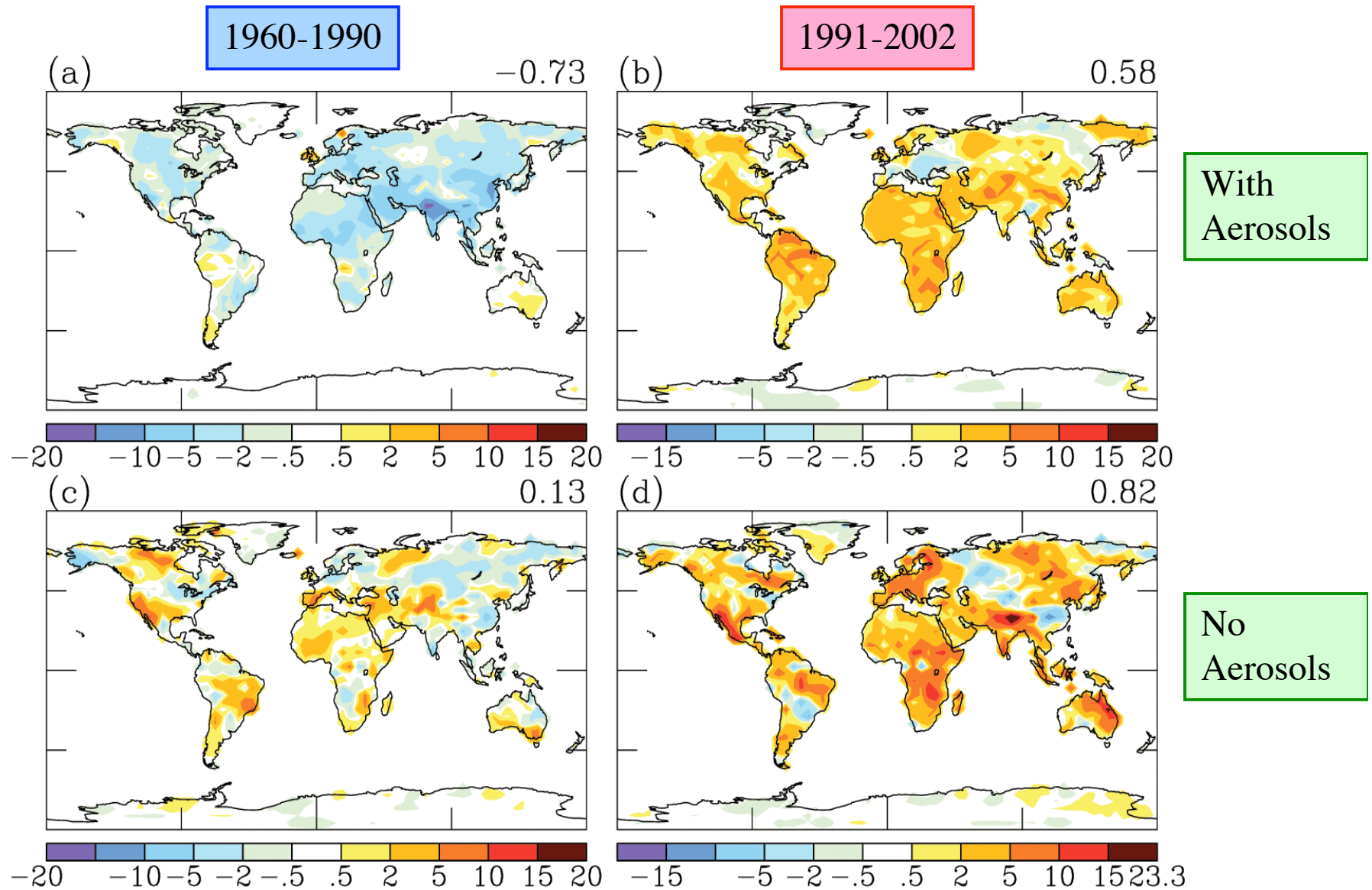
Exp B: Similar to Exp A but no anthropogenic aerosols



Units (W m^{-2}); Global means: r.h.s of graph

(Nazarenko and Menon 2005, GRL)

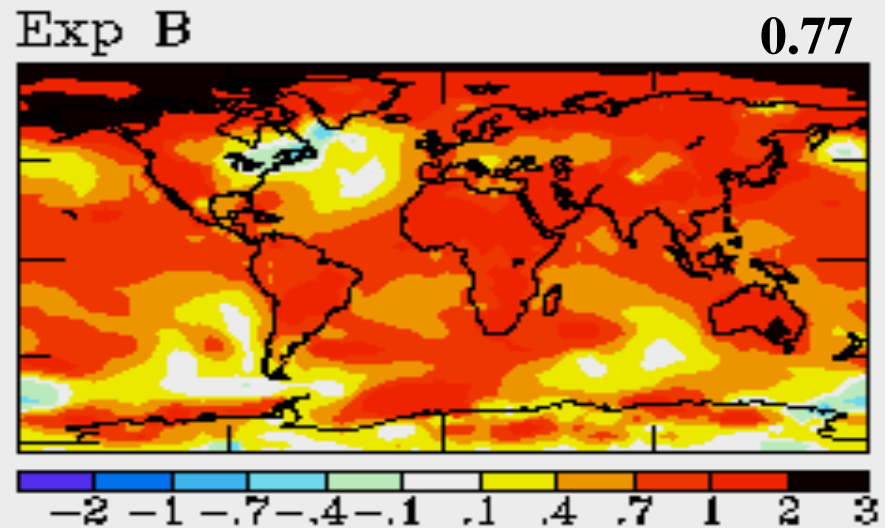
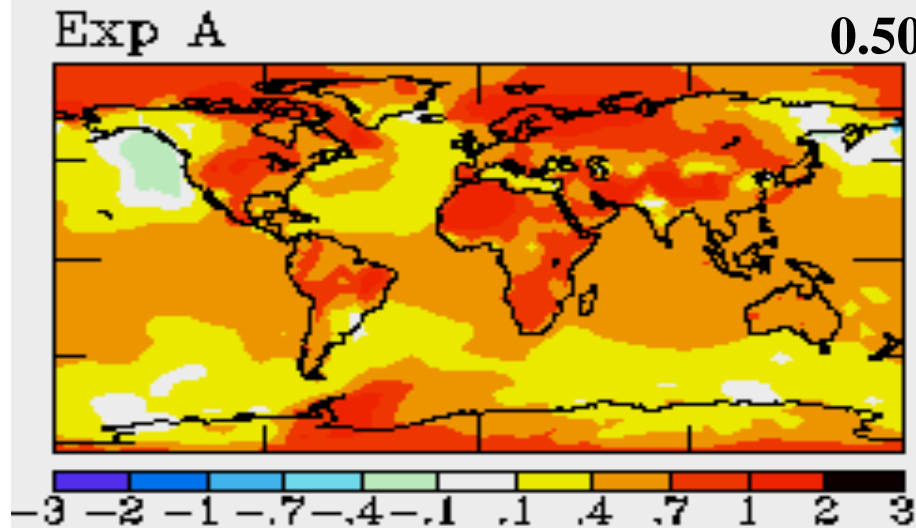
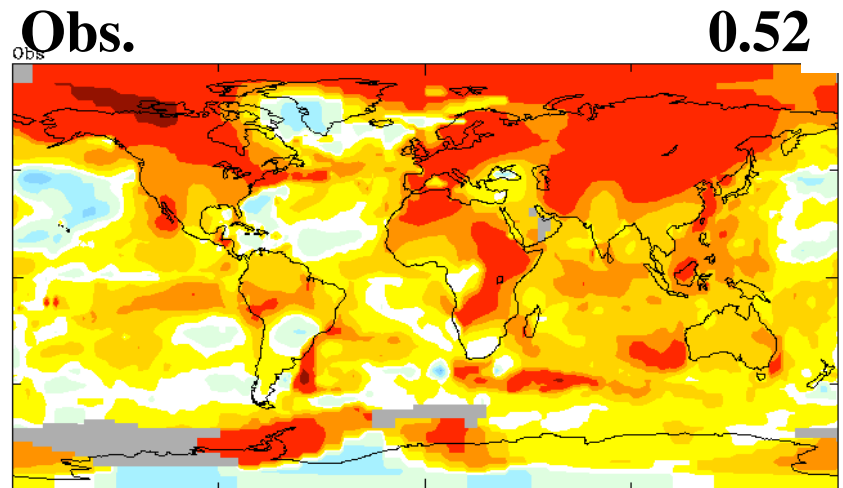
Trends in Global Surface Absorbed Solar Radiation



Surface Temperature Trends: 1960-2002

With anthropogenic aerosols, temperature trends (Exp A) match observed trends.

Policy Implications: Without mitigating GHGs, sfc. temp. reduction due to aerosols may no longer mask GHG effects in some regions if only aerosols are reduced.



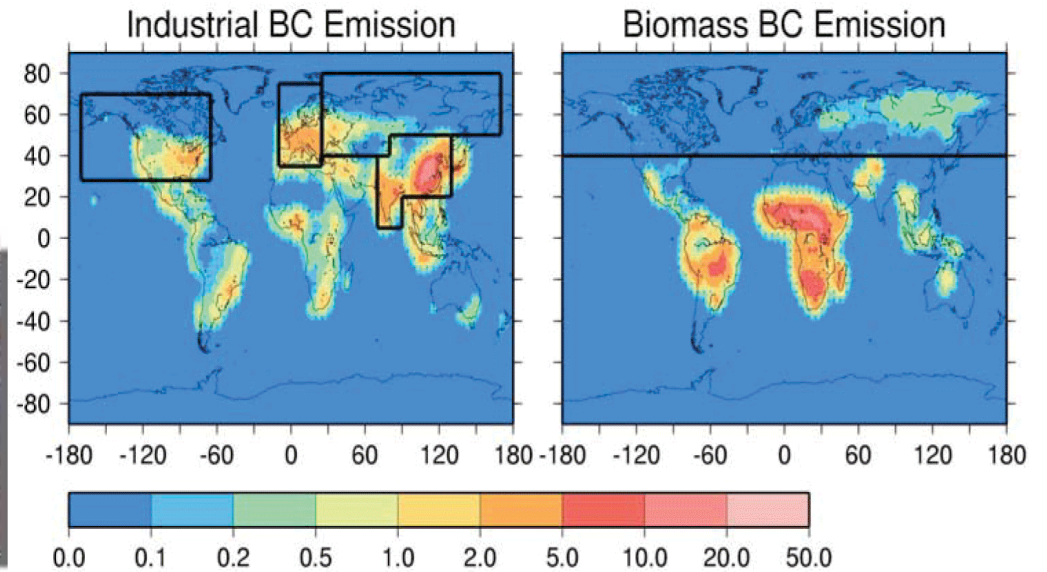
Units in K; Global means: r.h.s. of graph

Black Carbon Sources and Distribution

Black carbon is a product of **incomplete combustion.**



Black carbon Distribution



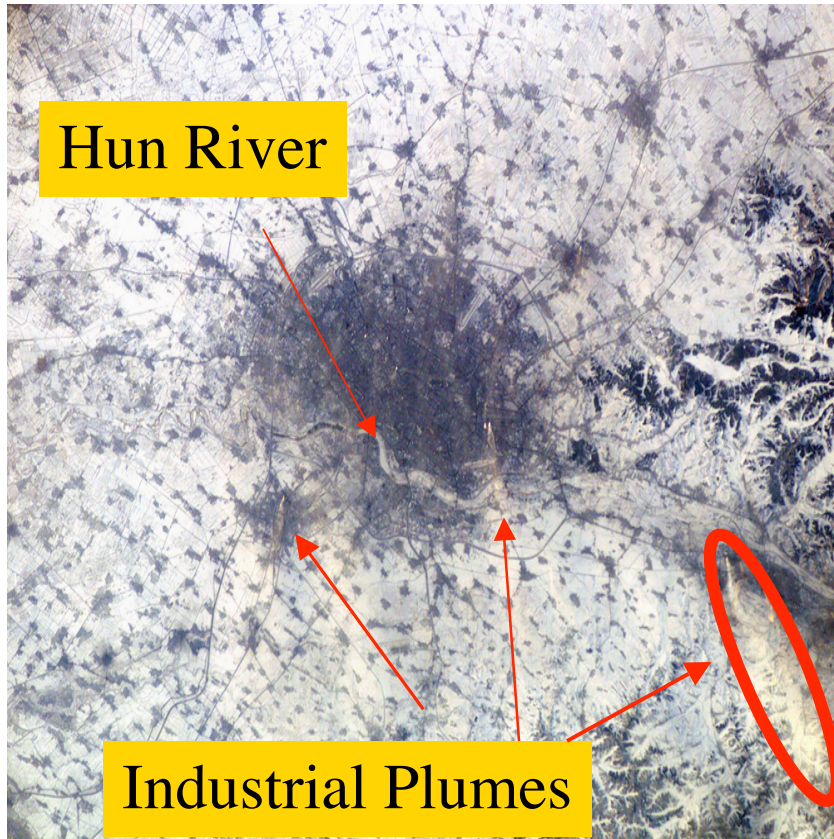
$(\text{ng C m}^{-2} \text{ s}^{-1})$

Images: <http://www.asthmacure.com>, NASA GSFC

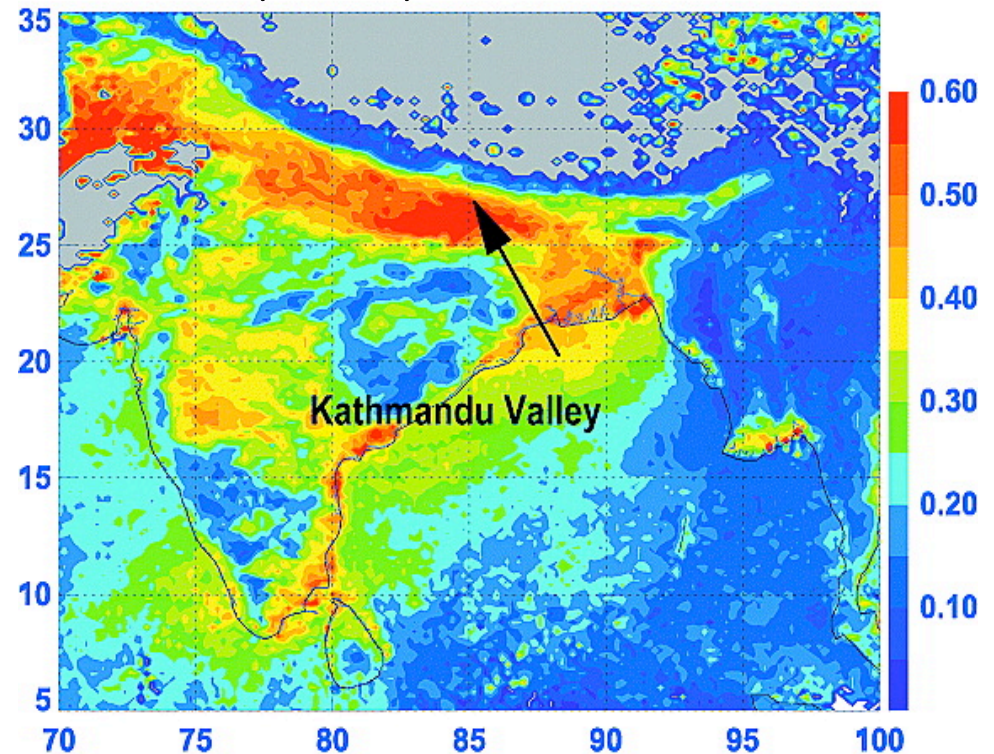
Koch et al. 2005

Industrial Pollution: Shenyang, China, and India/Nepal

Snow cover change between urban and rural areas



Aerosol optical depth from satellite (MODIS), Dec 2002



- Aerosol visible optical depths ~ 0.6
- Aerosol single scattering albedo ~ 0.78
- Inferred shortwave atmospheric forcing $\sim 25 \text{ W m}^{-2}$

(Ramana et al. 2004, GRL)

Image courtesy: Image Analysis Laboratory, NASA
Johnson Space Center

Astronaut photograph ISS010-E-13807, acquired
January 18, 2005

China: Black Carbon and Summer Monsoon Trends

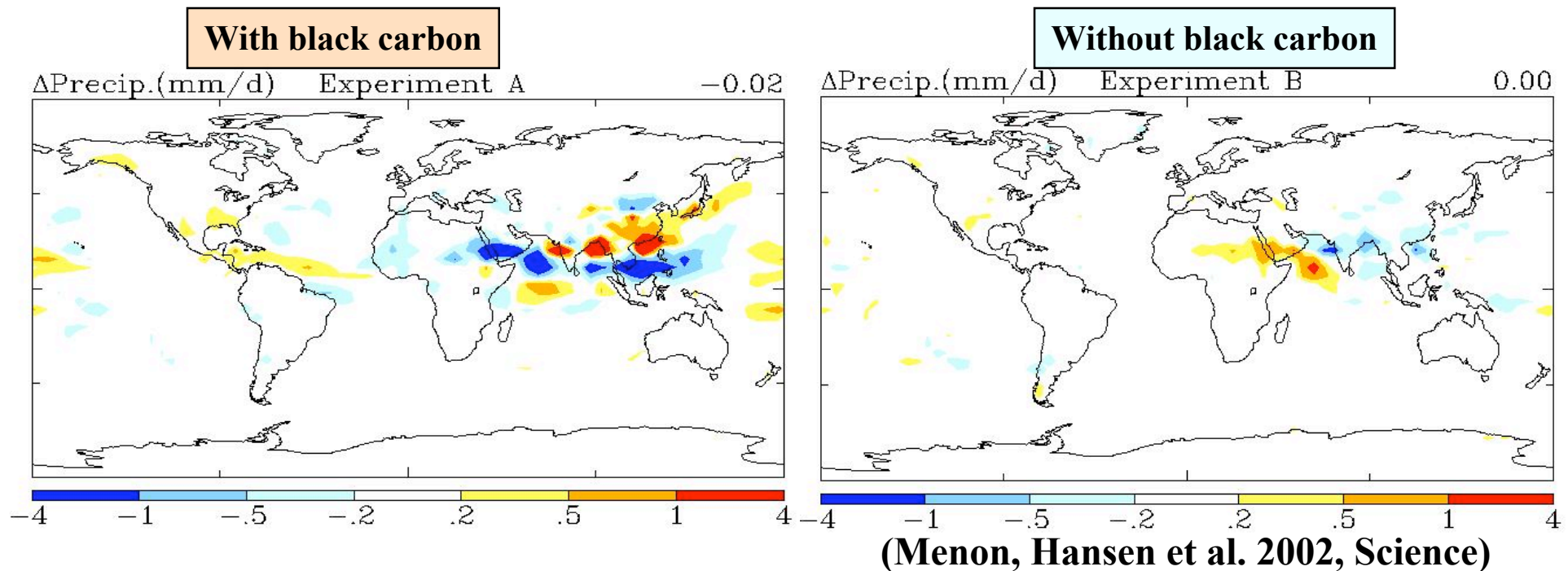
Observations over the last few decades:

- increased **floods**/droughts in the **south**/north; increased dust storms in the spring;
- precipitation trends **largest observed since 950 AD**.

We link increased emissions over China (since the late 1970's) with observed climate.

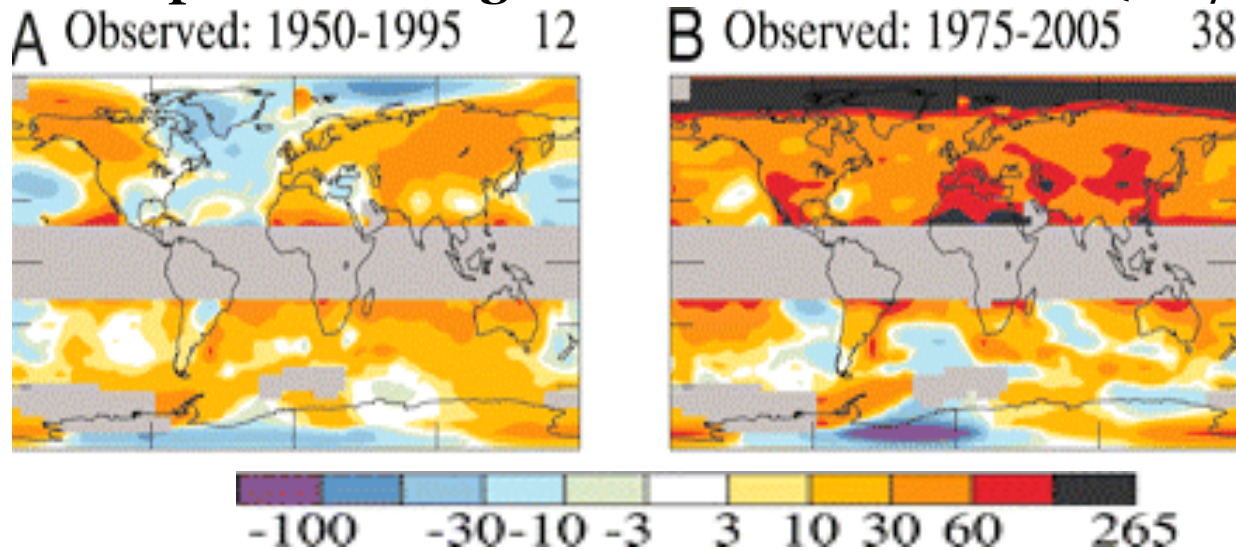
Assume a large proportion of aerosols are absorbing (**black carbon**).

Changes in **heating profile** affects convection, stability and spatial distribution of precipitation.



Climate impacts on species

Observed poleward migration rate of isotherms (km/decade)



Based on a study of 1700 species: Poleward migration of 6 km/decade and vertical migration in alpine regions of 6 m per decade in the second half of the 20th century (Parmesan and Yohe, 2003, Nature)

Global warming of **3°C** over the 21st century could eliminate **~60%** of species.

In the past several mass extinctions of **50–90%** of species for global temperature changes of **~5°C**. (Hansen et al. 2006, PNAS)

Climate and extreme weather

1930's US Dust Bowl



Cooler than normal tropical Pacific Ocean surface temperatures and warmer tropical Atlantic Ocean temperatures created shifts in the large-scale weather patterns and low level winds that reduced the normal supply of moisture from the Gulf of Mexico and inhibited rainfall throughout the Great Plains. (Schubert et al., 2004, Science)

Satellite image of hurricane



Sea-surface temperatures have a strong influence on hurricanes.

For 1906–2005: 84% chance that **external forcing** explains at least 67% of **observed SST increases** in the two tropical cyclogenesis regions (Santer et al., 2006, PNAS)

OUTLINE

Climate change - ~1 M yrs to present-day

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Future Climate

Arctic sea-ice decline

1982



2007



Melting permafrost peatlands (Noyabrsk, Western Siberia)

SI Snow and Ice Data Center, 2007



2010 - 2030



2070 - 2090



Climate Impact Assessment, 2004

Arctic sea-ice decline

1982



2007



National Snow and Ice Data Center, 2007



2010 - 2030



2040 - 2060



2070 - 2090

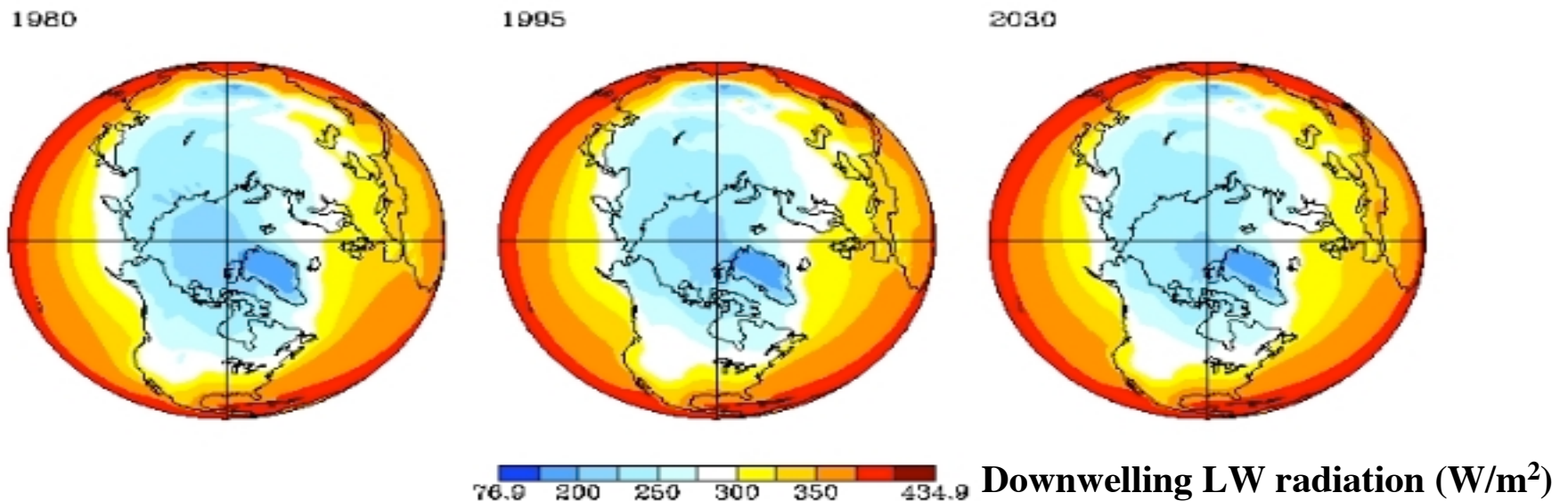
Arctic Climate Impact Assessment, 2004

Aerosol Impacts on the Arctic

For short-term mitigation, we have investigated future climate change for short-lived species (ozone and aerosols).

Our modeling work indicates that aerosol effects on cloud properties can affect Arctic surface heating more than ozone does (Menon et al. 2008, ERL).

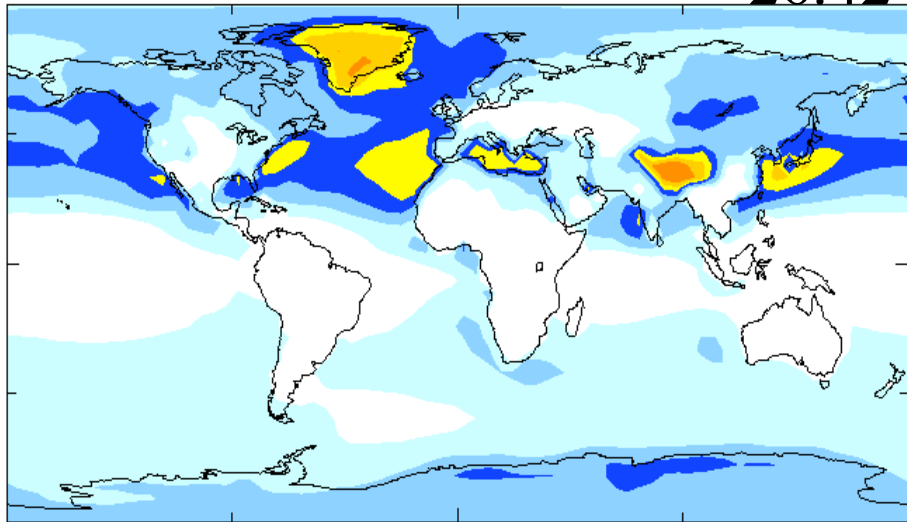
For the future scenario (2030) considered, increased downwelling longwave (LW) radiation, due to cloud changes, implies increased surface warming.



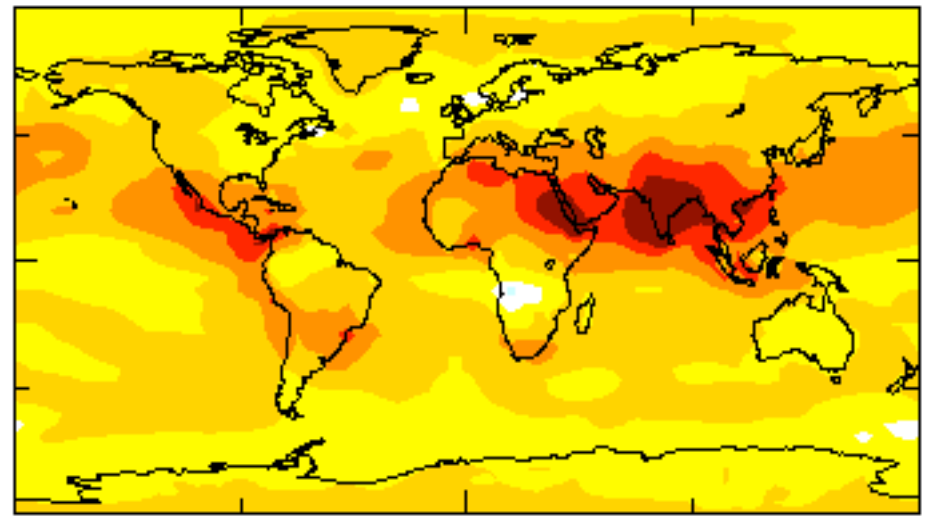
Future Air Quality

Global annual amount and change in surface amount

1995 Surface Ozone 26.42



Change in Surface Ozone 3.43

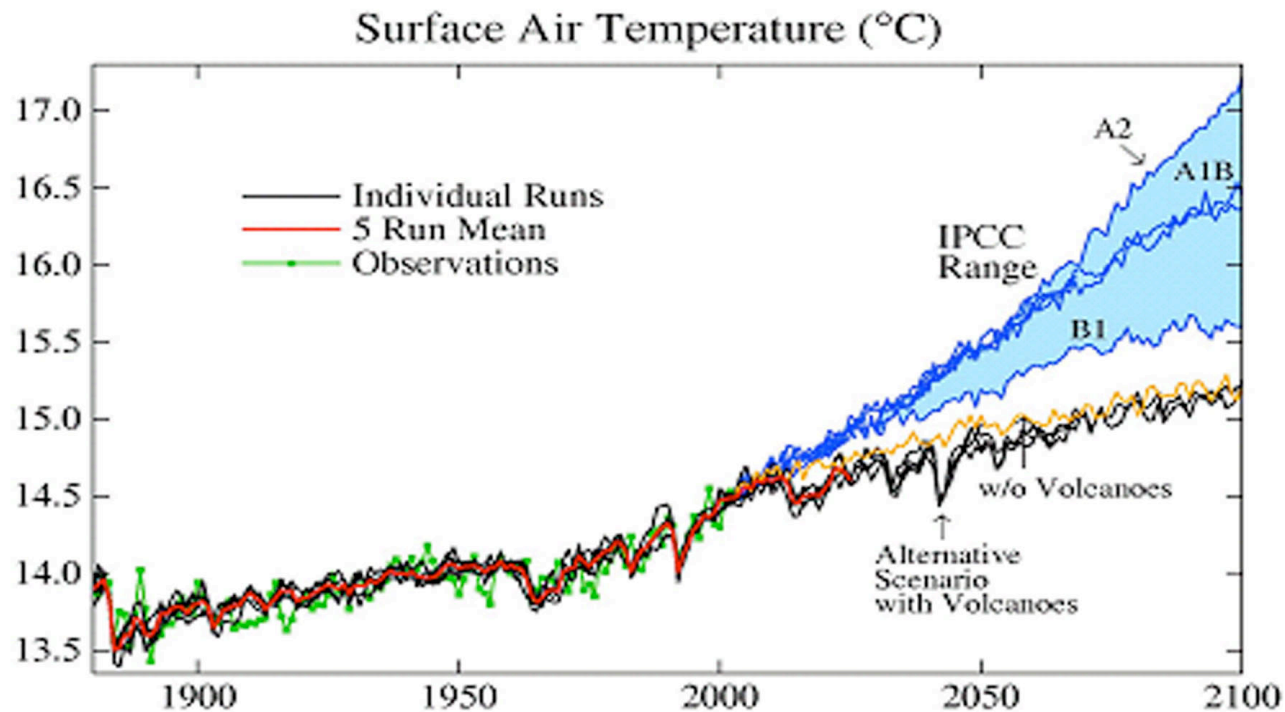


0 25 30 35 40 45 50 57.5 (ppbv) -10 -8 -6 -4 -2 -1 .5 2 4 8 12 30

Background ozone level of ~ 30 ppbv is exceeded in several tropical locations for the future scenario (2030) considered.

(Menon et al. 2008, ERL)

21st Century Global Warming

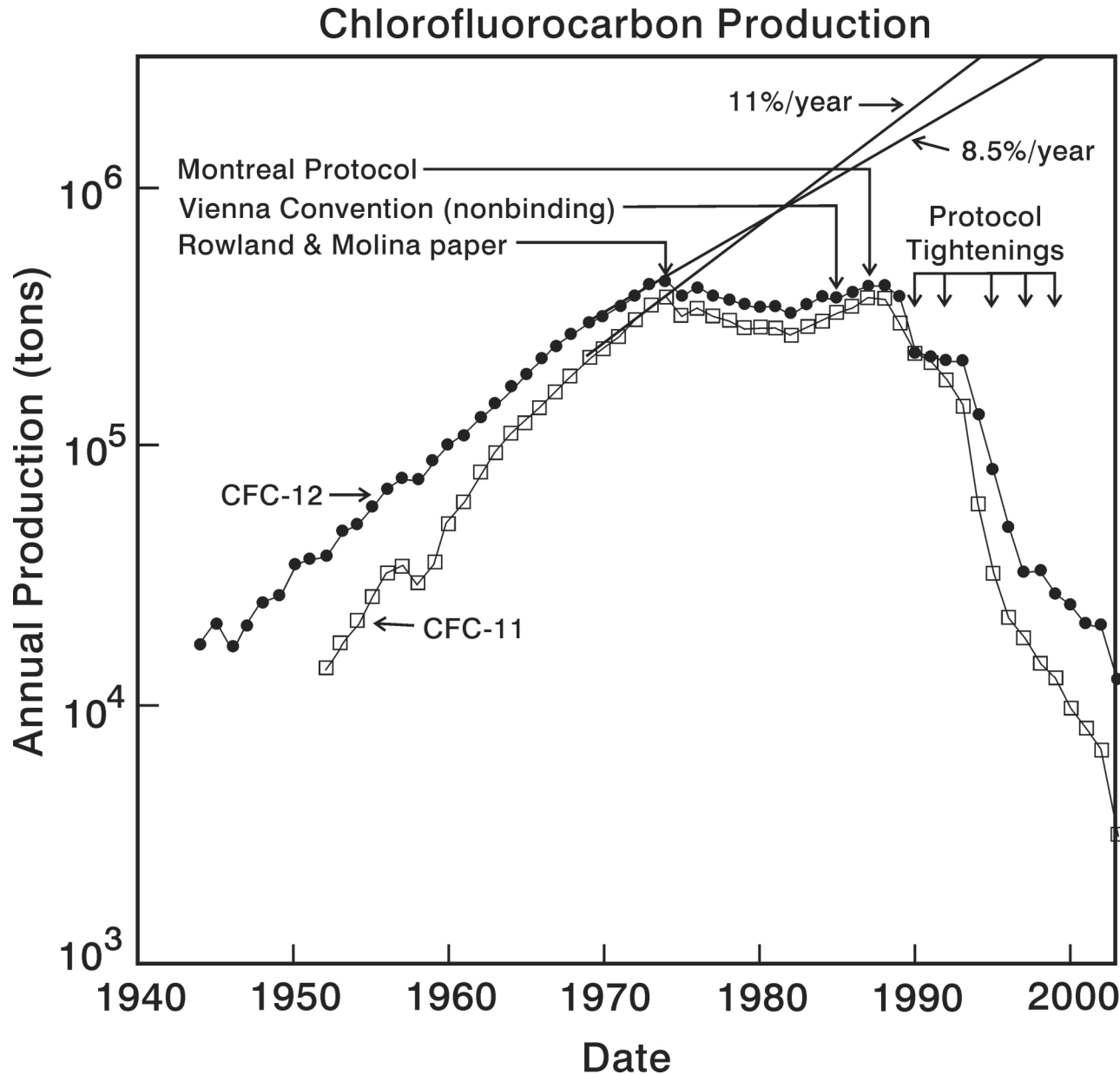


Climate Simulations for IPCC 2007 Report

- ▶ **Climate Model Sensitivity 2.7-2.9°C for 2xCO₂**
(consistent with paleoclimate data & other models)
- ▶ **Simulations Consistent with 1880-2003 Observations**
(key test = ocean heat storage)

Source: Hansen et al., 2005, 2007a,b.

CFC success story



Source: J Hansen

Climate Change Summary

Climate change: 1880 to Present

- Increased surface temperatures - from greenhouse gases.
- Global dimming: 1960-1990 and “dimming” reversal beyond 1990 observed only if anthropogenic aerosol effects are included.

Man-made aerosols directly impact climate and may mask GHG effects.

Future climate change: 2000 to 2100

- Depends on emission pathways followed that are guided by policy.
- Arctic region vulnerable to climate change from aerosols.
- Ozone and aerosol effects stronger over tropical regions such as India, where such effects were not dominant before.

Large unknown in climate change: emission sources and interactions/feedbacks of the climate system.

Acknowledgement

Funding Support:

NASA : Radiation Sciences Program

Modeling, Analysis and Prediction Program

LBNL: LDRD

DOE: Atmospheric Radiation Program,

Integrated Climate Assessment Program

Collaborators:

NASA Goddard Institute for Space Studies

Jim Hansen, Anthony Del Genio, Dorothy Koch, Nadine Unger, Larissa

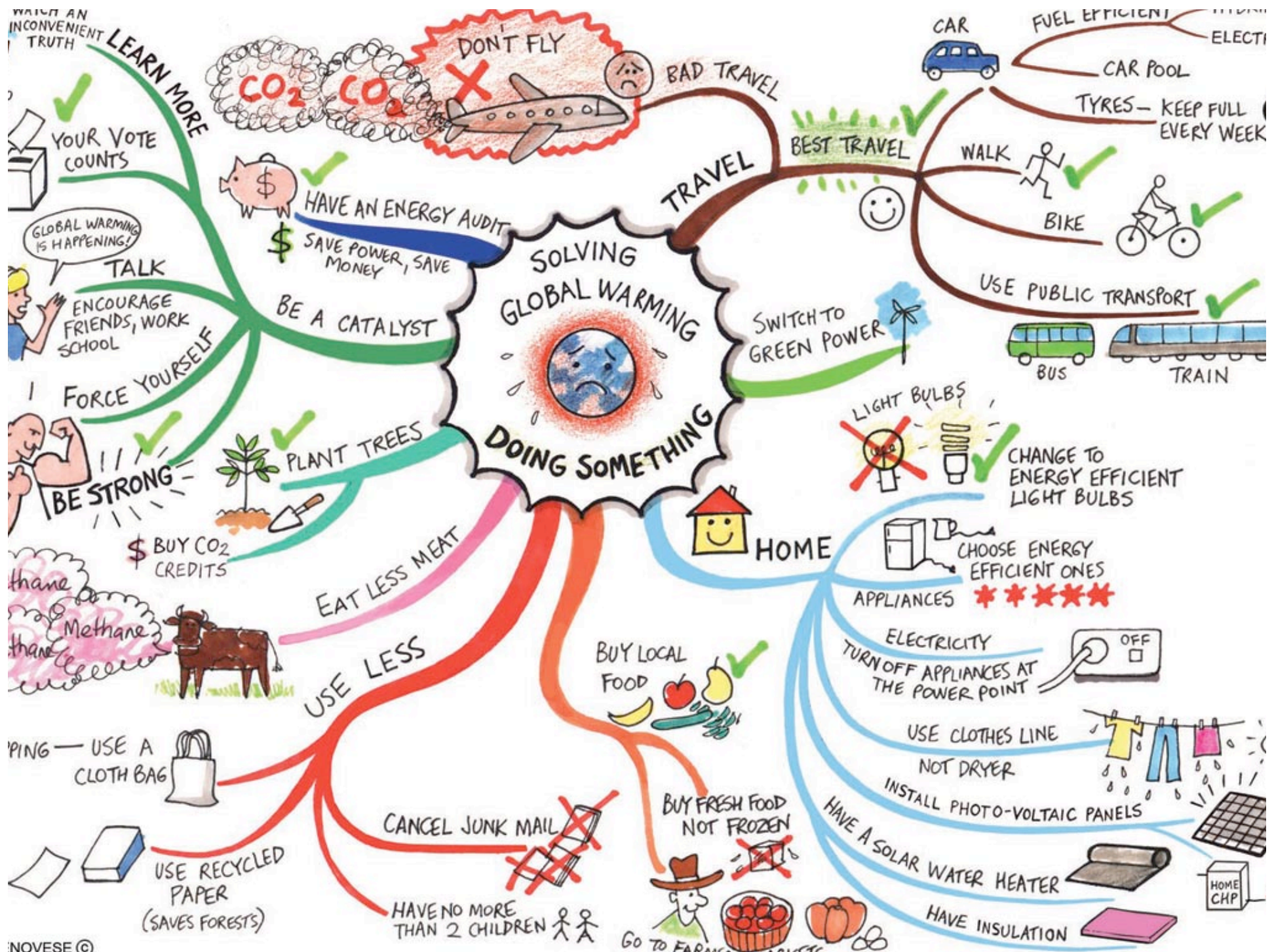
Nazarenko, Susanne Bauer

LBNL: Environmental Energy Technologies Division

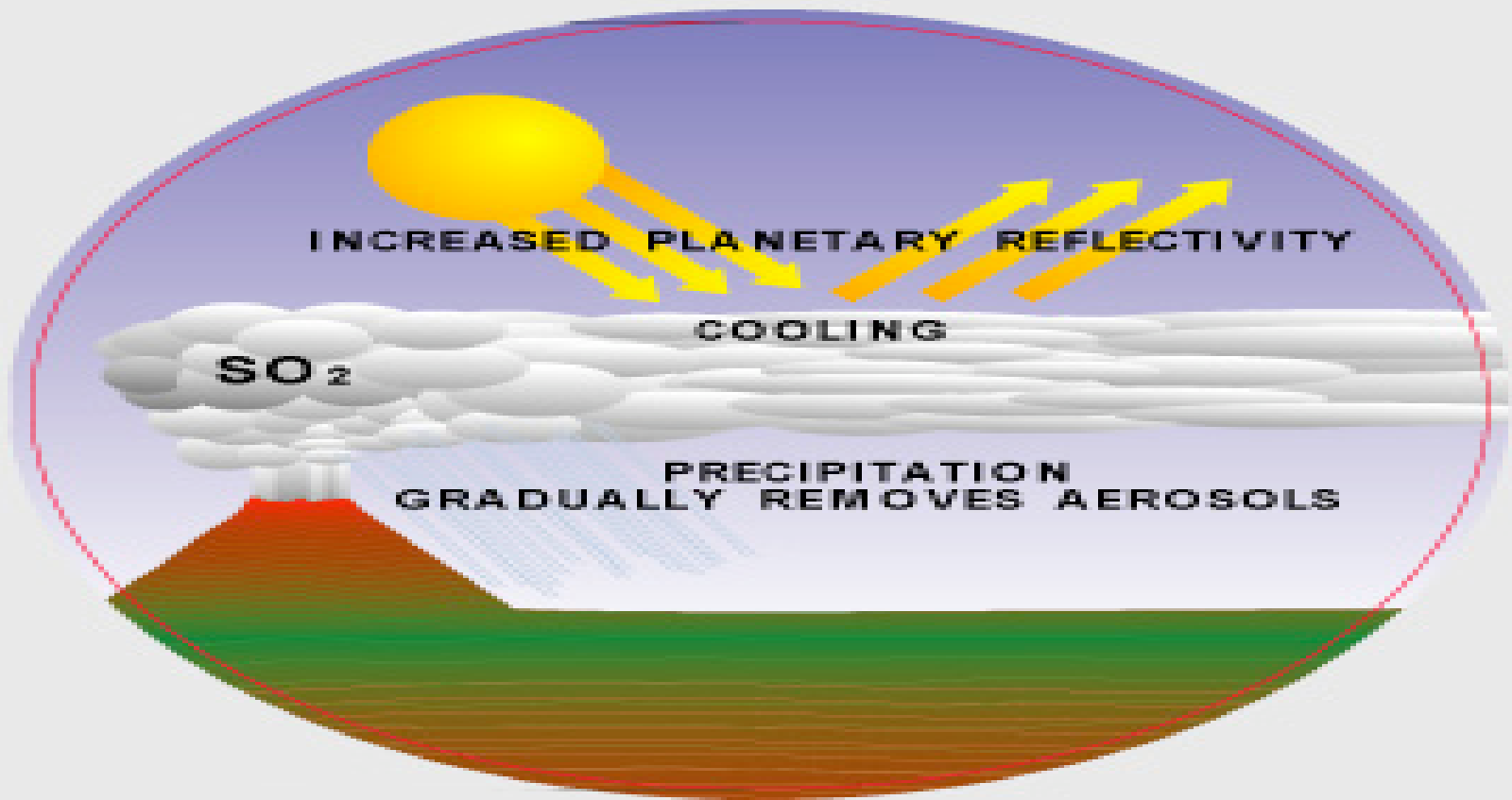
Tica Novakov, Igor Sednev

Summer students: Nick Burns, Nick Chmura, Parmendar Singh, Banglun Zheng,

Geeta Persad

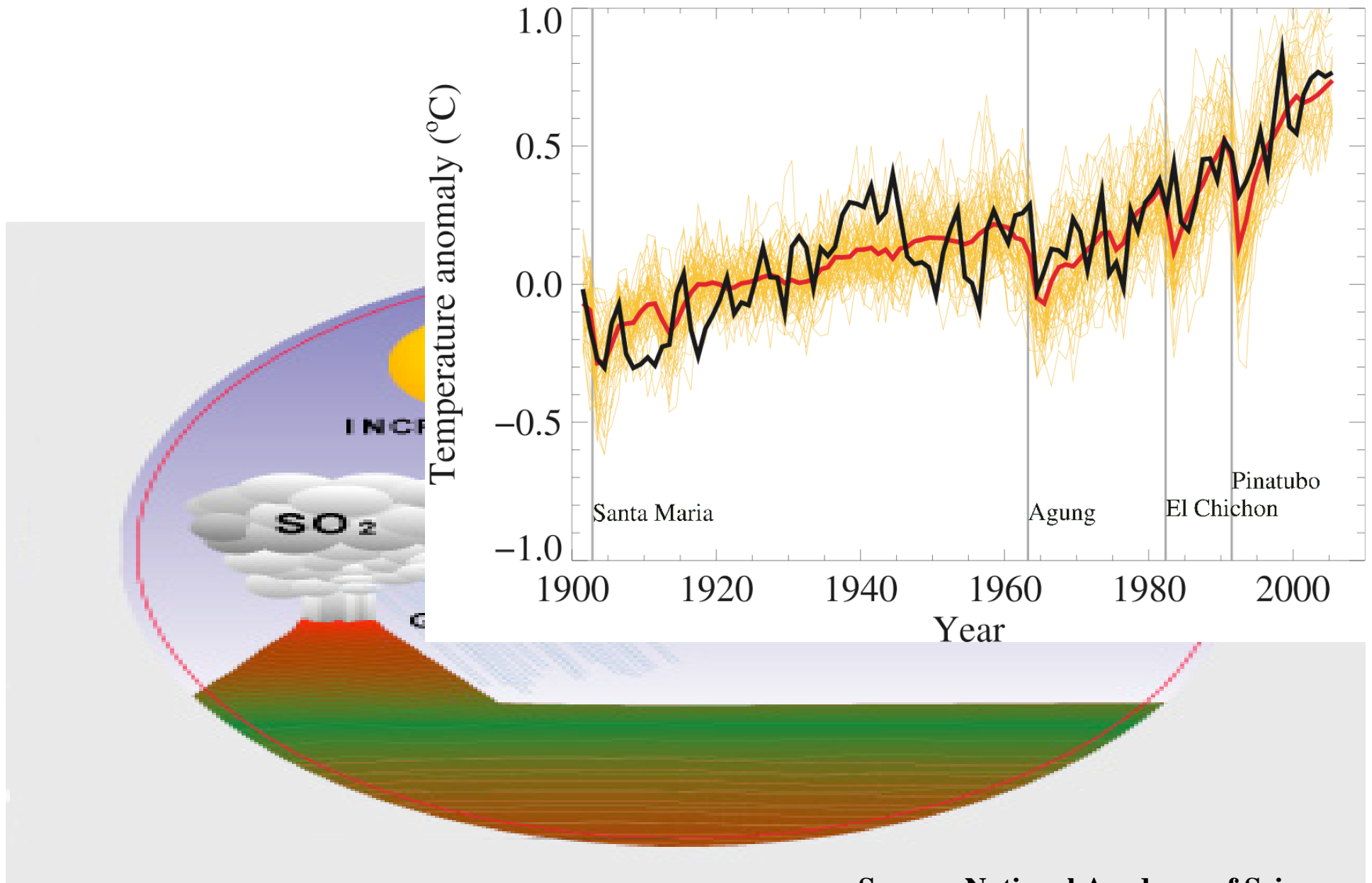


Geo-engineering Climate



Source: National Academy of Sciences

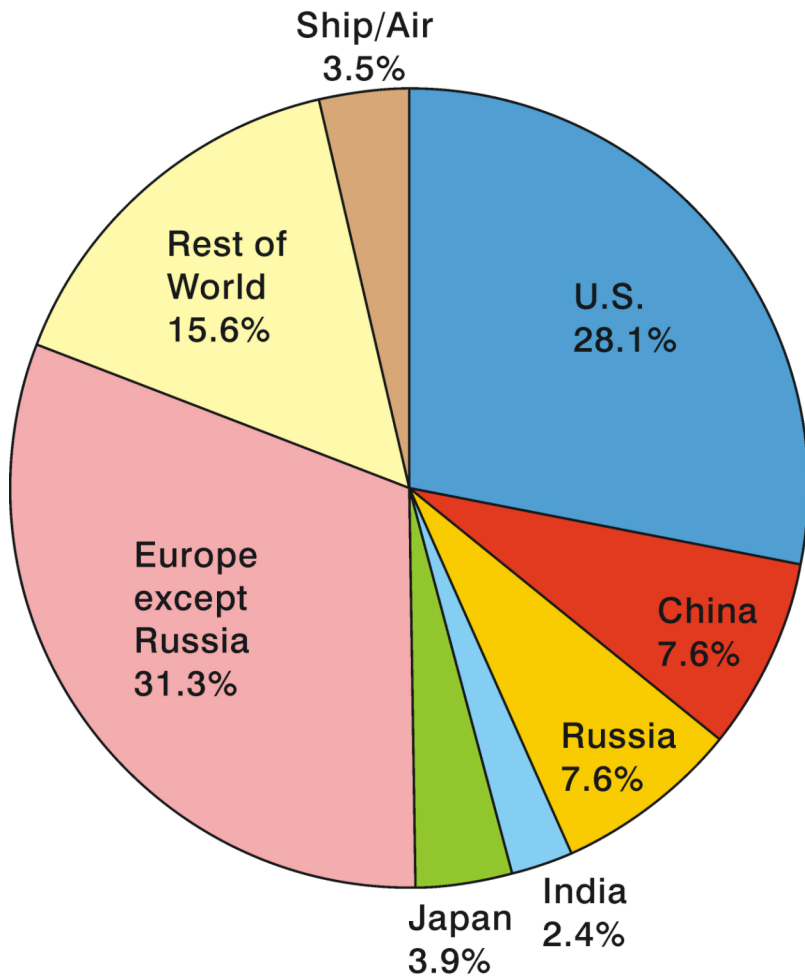
Geo-engineering Climate



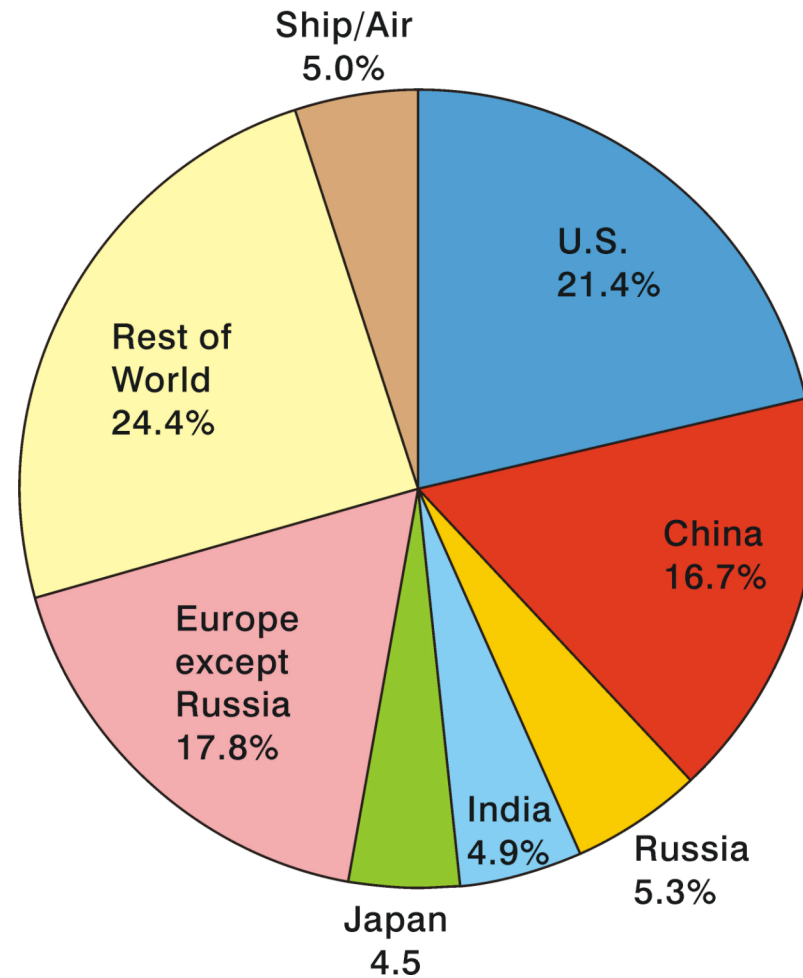
Source: National Academy of Sciences

Fossil Fuel CO₂ Emissions

Accumulated Fossil Fuel CO₂ (1850-2004)



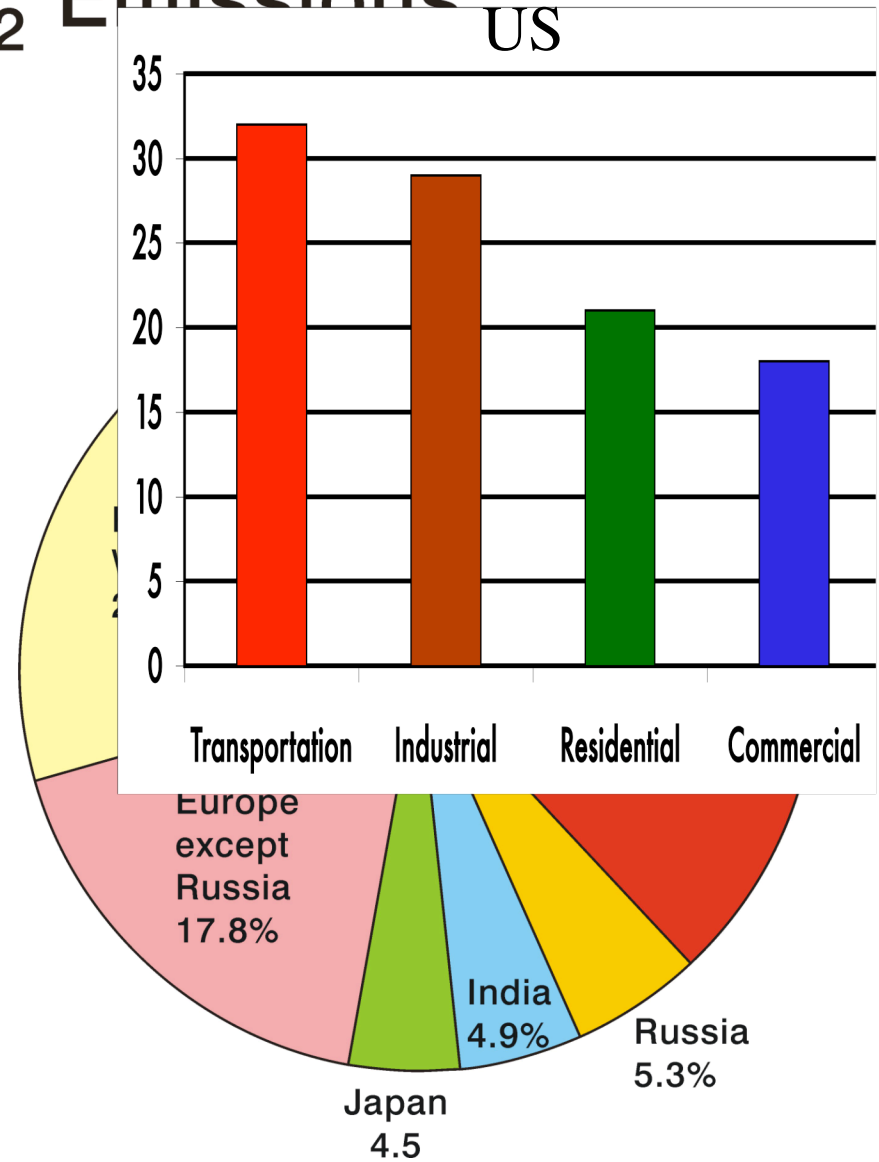
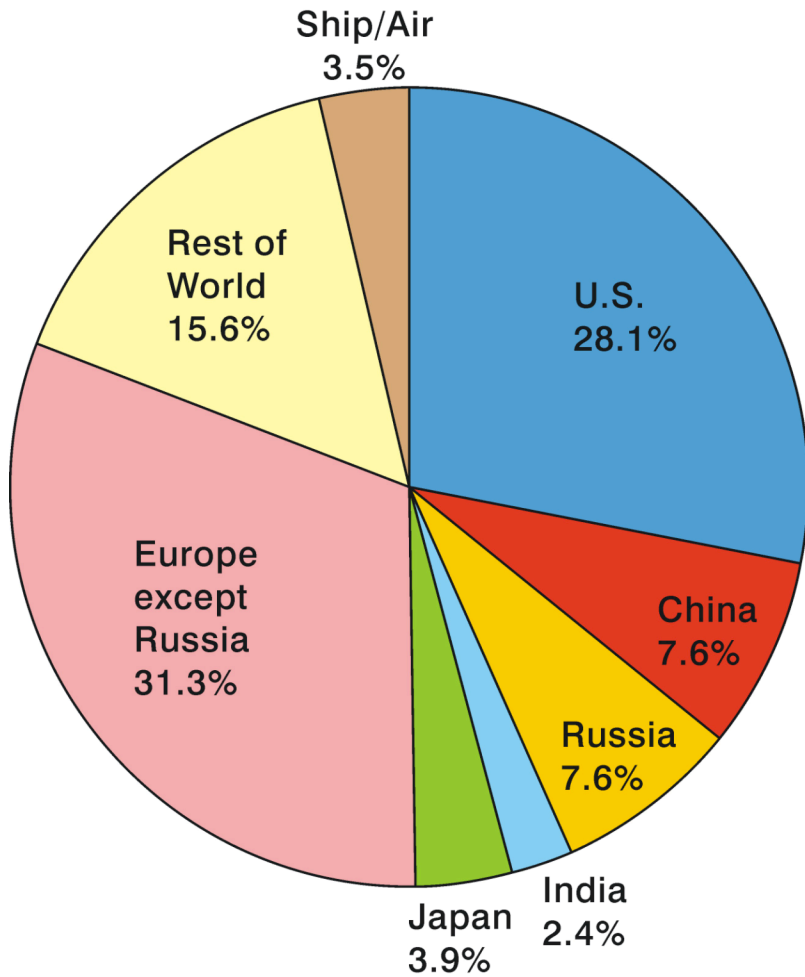
2004 Portions of CO₂ Emissions



Source: J Hansen

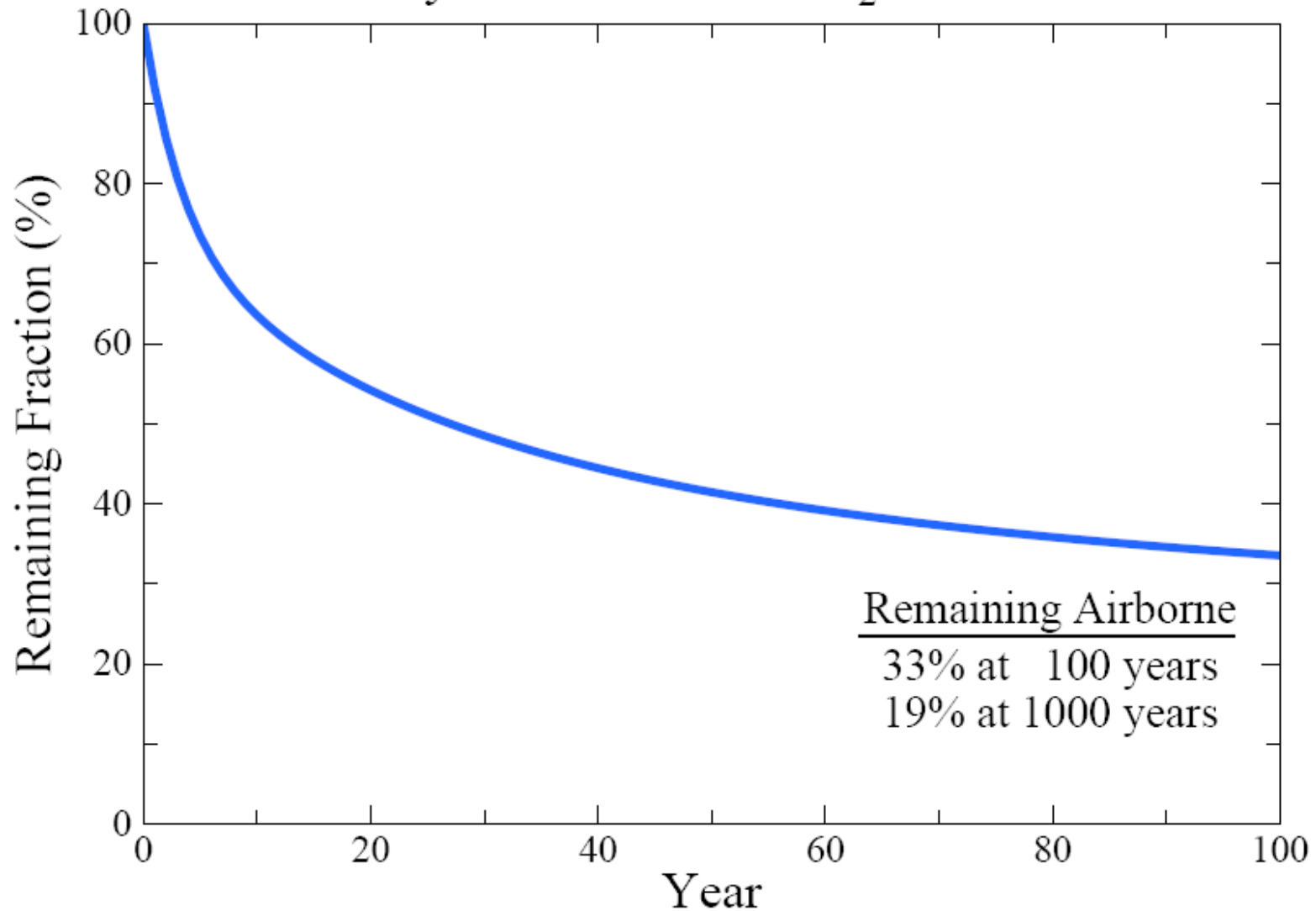
Fossil Fuel CO₂ Emissions US

Accumulated Fossil Fuel CO₂ (1850-2004)



Source: J Hansen

Decay of Fossil Fuel CO₂ Emission



The fraction of CO₂ remaining in the air, after emission by fossil fuel burning, declines rapidly at first, but 1/3 remains in the air after a century and 1/5 after a millennium (Hansen et al. 2007, *Atmos. Chem. Phys.* **7**, 2287-2312, 2007).

Methane more potent than CO₂

GWP: Ratio of heat trapped by one unit mass of the GHG to that of one unit mass of CO₂ over a specified time period.

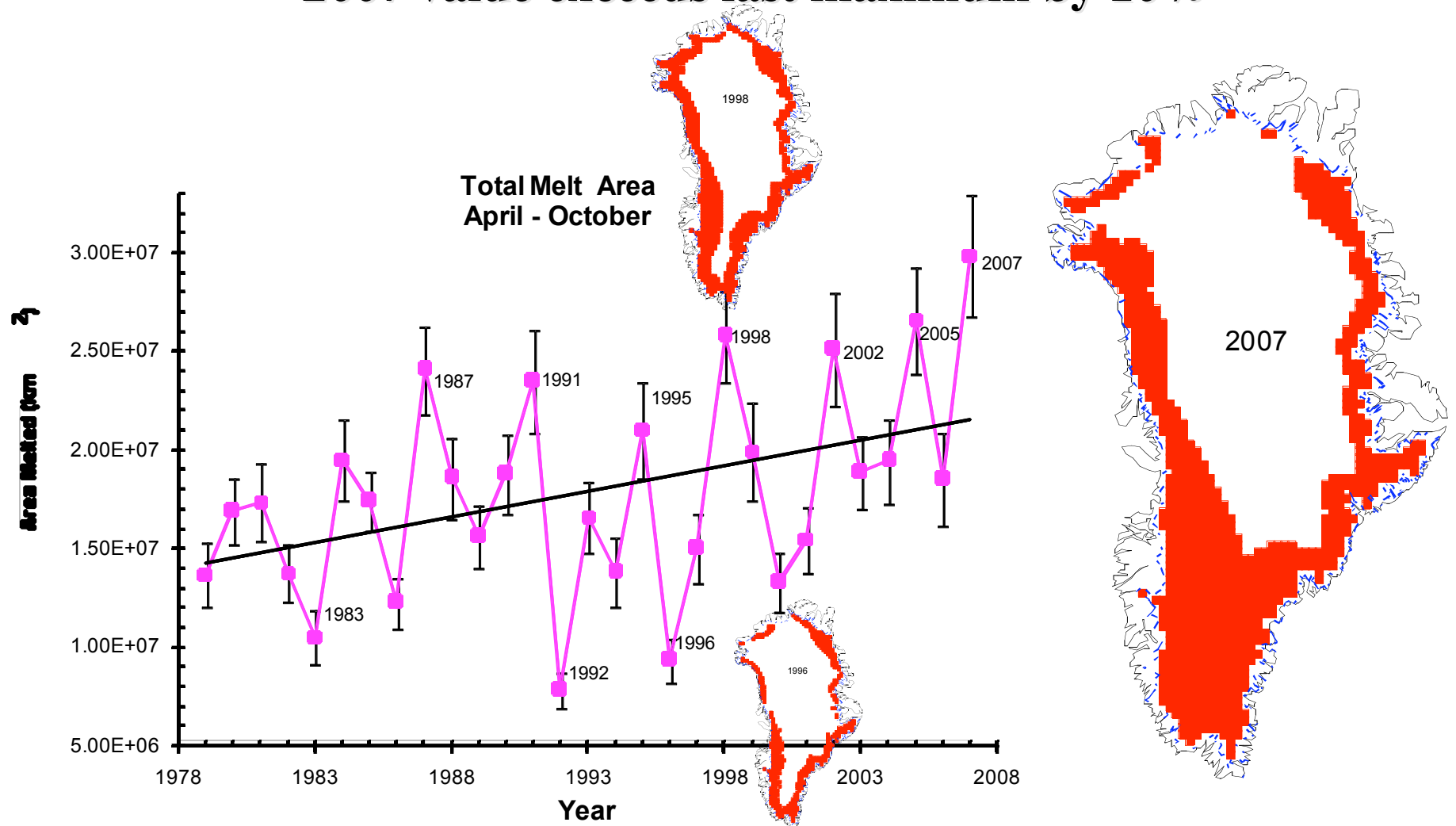
As the permafrost thaws as a result of global warming caused by increasing CO₂, large quantities of CH₄ are released-- causes more warming and the release of even more CH₄.

Effects of methane

- Helps control amount of OH (hydroxyl) in the troposphere
- Affects concentrations of water vapor and O₃ in the stratosphere
- Reaction with hydroxyl (OH) radical (~90%) in the troposphere
 - OH is formed by photodissociation of tropospheric ozone and water vapor
- OH is the primary oxidant for most tropospheric pollutants (CH₄, CO, NO_x)
- Amount CH₄ removed constrained by OH levels and reaction rate
- Dissociation of CH₄ leads to CO₂: additional climatic forcing

Greenland Total Melt Area

2007 value exceeds last maximum by 10%



Source: J Hansen

Konrad Steffen and Russell Huff, CIRES, University of Colorado at Boulder

Ice Age Climate Forcings (W/m^2)

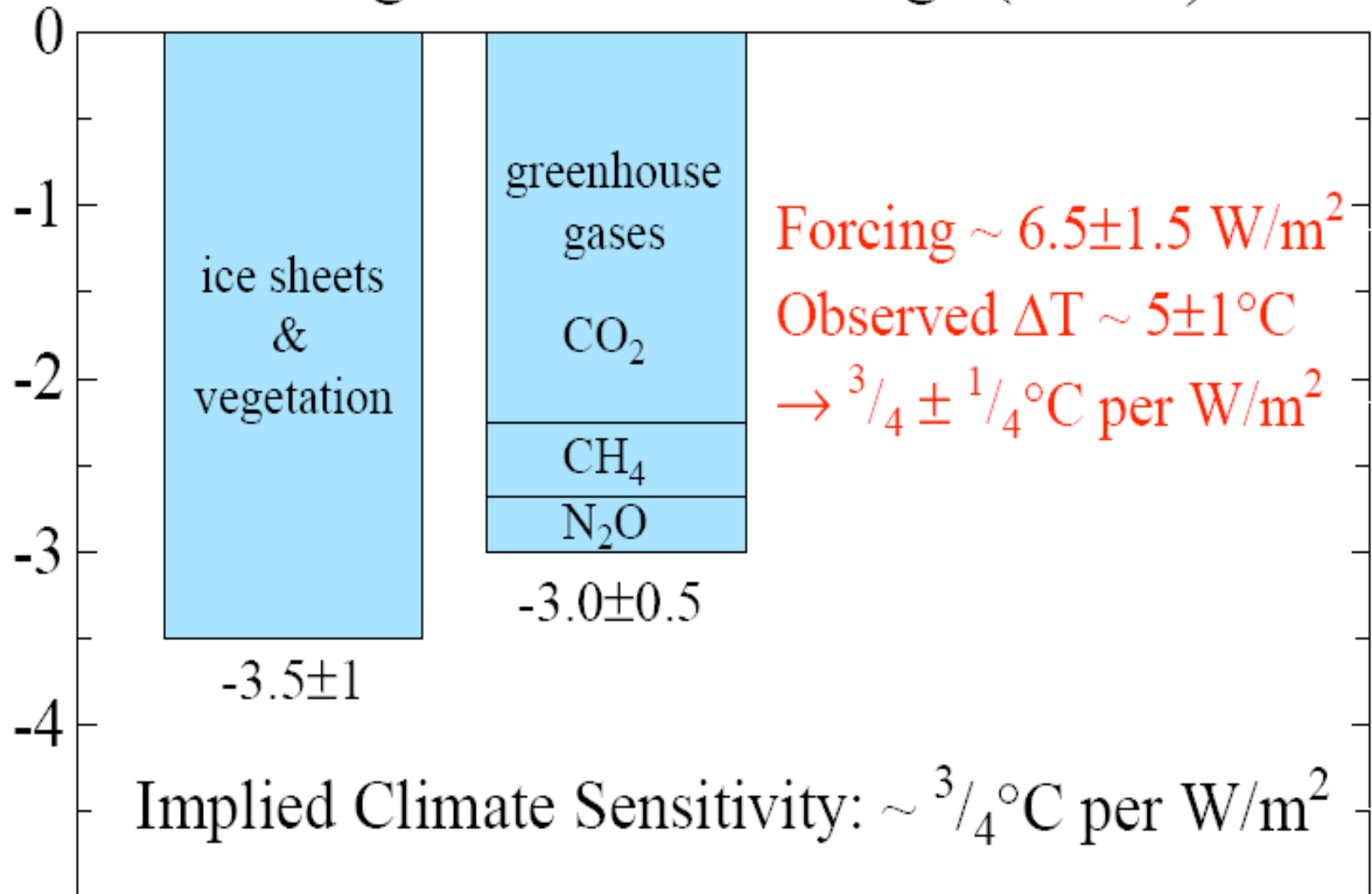


Fig. S2. Climate forcings during ice age 20 ky BP, relative to the present (pre-industrial) interglacial period.

Source: Hansen, J

Assessment of Target CO₂

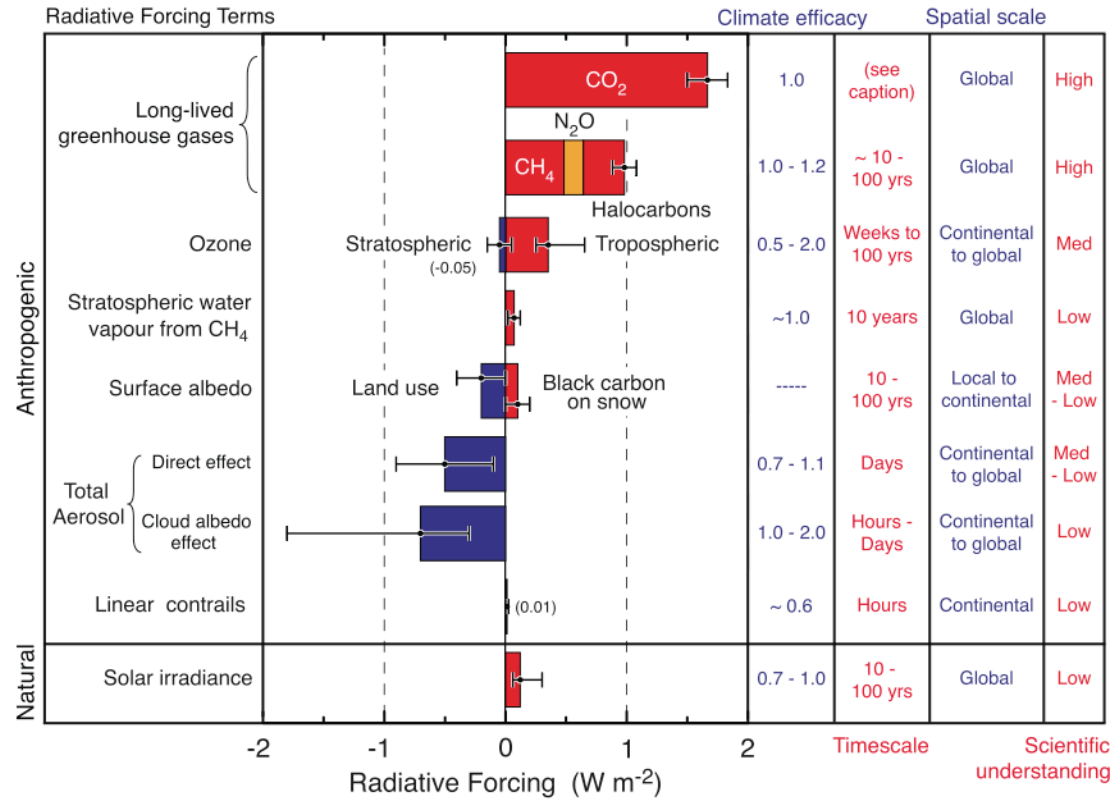
<u>Phenomenon</u>	<u>Target CO₂ (ppm)</u>
1. Arctic Sea Ice	300-325
2. Ice Sheets/Sea Level	300-350
3. Shifting Climatic Zones	300-350
4. Alpine Water Supplies	300-350
5. Avoid Ocean Acidification	300-350

◇ Initial Target CO₂ = 350* ppm

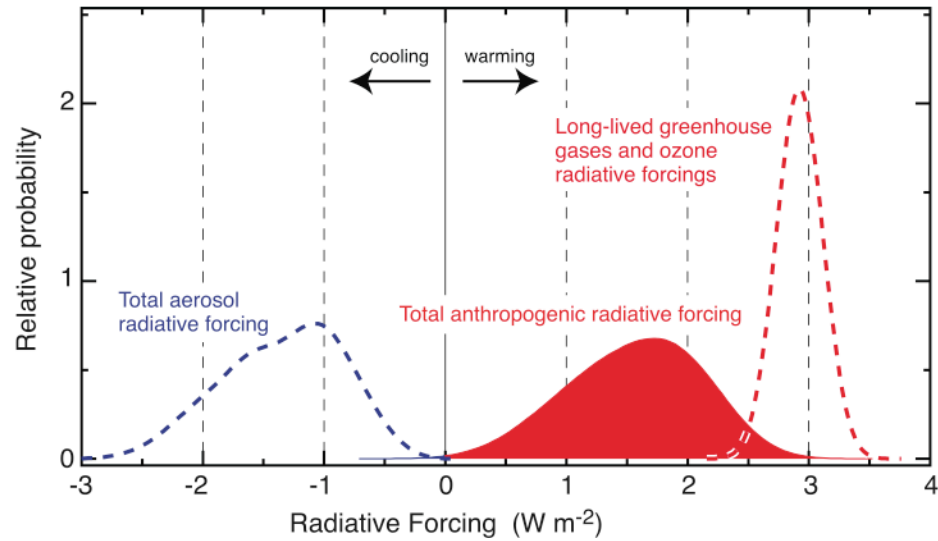
*assumes CH₄, O₃, Black Soot decrease

Radiative forcing of climate between 1750 and 2005

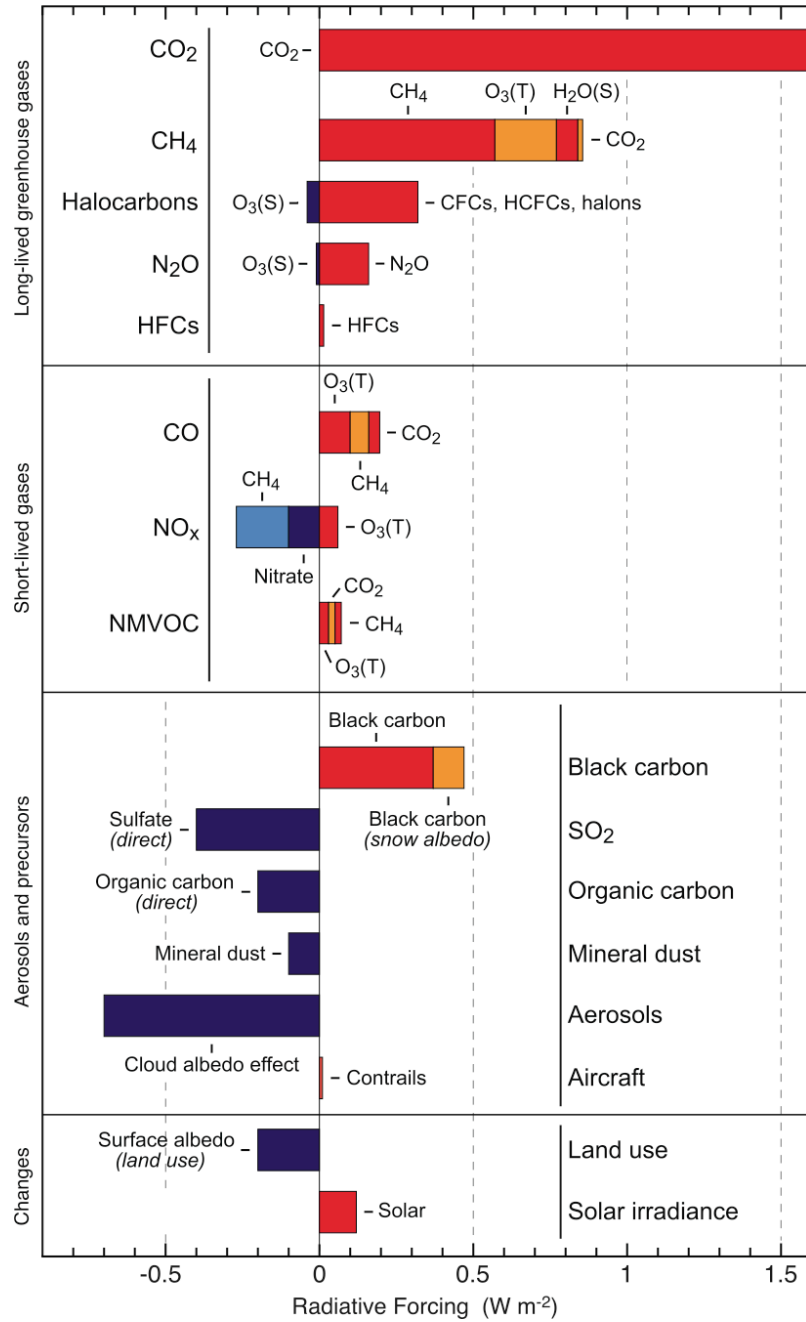
A.



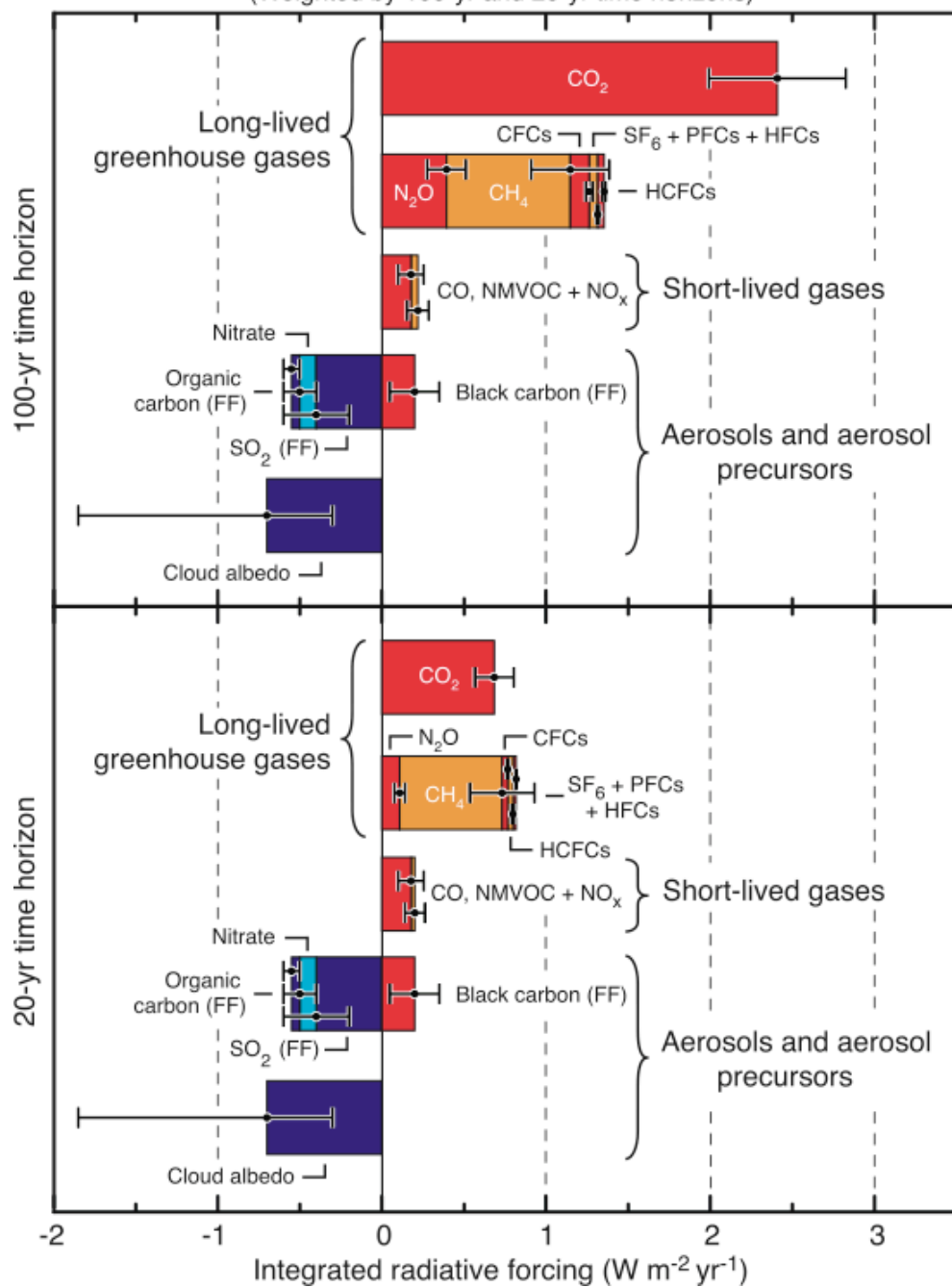
B.



Components of radiative forcing for principal emissions



Integrated Radiative Forcing for Year 2000 Global Emissions
(Weighted by 100-yr and 20-yr time horizons)



Pliocene: 5.332 million to 1.806 million years before present

Zanclean (5.3 - 3.6 Ma)

Piacenzian (3.6 - 2.6 Ma)

Gelasian (2.6 - 1.8 Ma)

Pleistocene: 1.8 million to 10,000 years BP covering recent period of repeated glaciations

Early Pleistocene (1.8 - 0.78 Ma)

Middle Pleistocene (780 - 130 ka)

Late Pleistocene (130 - 10 ka)

Older Dryas (14 - 13.6 ka)

Aller'd (13.6 - 12.9 ka)

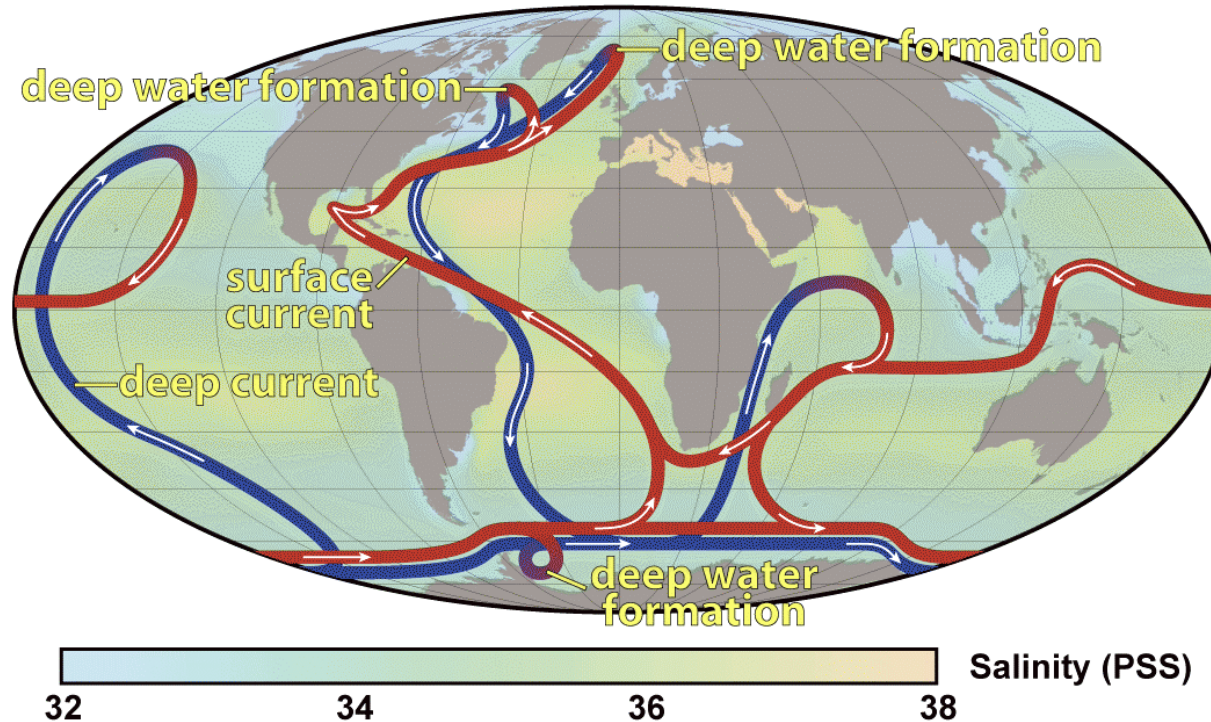
Younger Dryas (12.9 - 11.5 ka) Brief ($\sim 1300 \pm 70$ years) cold climate period approximately 12,800 to 11,500 years BP.

Holocene (10 ka - present)

Boreal and Atlantic

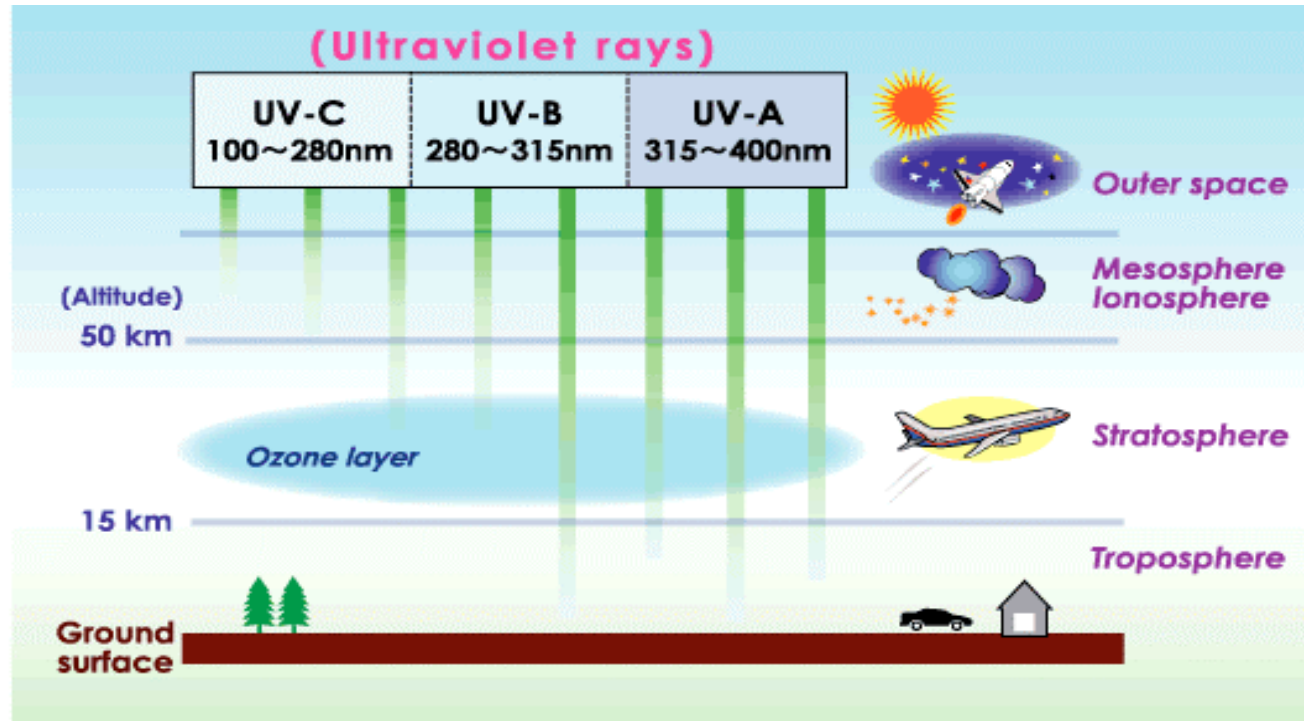
- The prevailing theory holds that the Younger Dryas was caused by a significant reduction or shutdown of the North Atlantic thermohaline circulation in response to a sudden influx of fresh water from Lake Agassiz and deglaciation in North America

Thermohaline Circulation



Large-scale ocean circulation that is thought to be driven by global density gradients created by surface heat and freshwater fluxes.

Global warming could, via a shutdown or slowdown of the thermohaline circulation, trigger localized cooling in the North Atlantic and lead to cooling, or lesser warming, in that region. This would affect in particular areas like [Iceland](#), [Ireland](#), the [Nordic countries](#), and [Britain](#) that are warmed by the [North Atlantic drift](#).



Each spring in the stratosphere over Antarctica, atmospheric ozone is rapidly destroyed by chemical processes. As winter arrives, a vortex of winds develops around the pole and isolates the polar stratosphere. When temperatures drop below -78°C , thin clouds form of ice, nitric acid, and sulphuric acid mixtures.

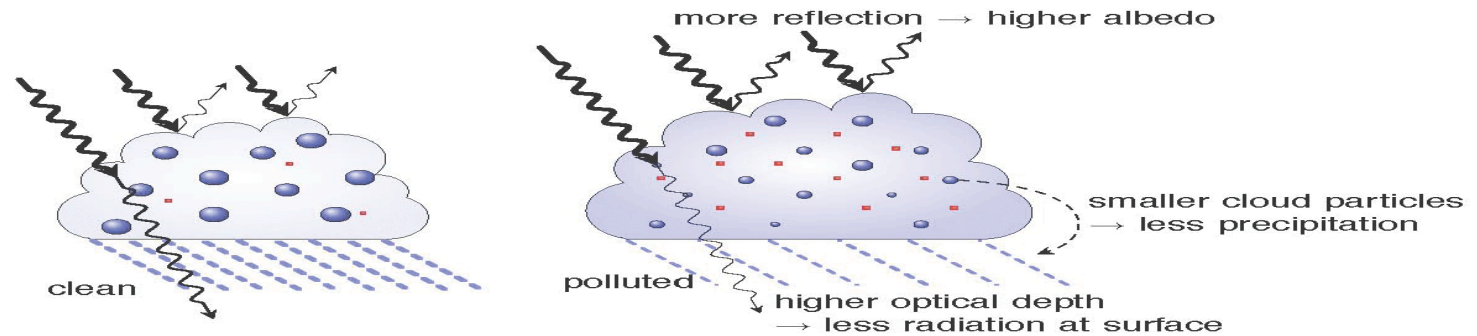
Chemical reactions on the surfaces of ice crystals in the clouds release active forms of CFCs. Ozone depletion begins, and the ozone hole appears.

Over the course of 2-3 months, $\sim 50\%$ of the total column ozone disappears. At some levels, the losses approach 90%. This has come to be called the Antarctic ozone hole.

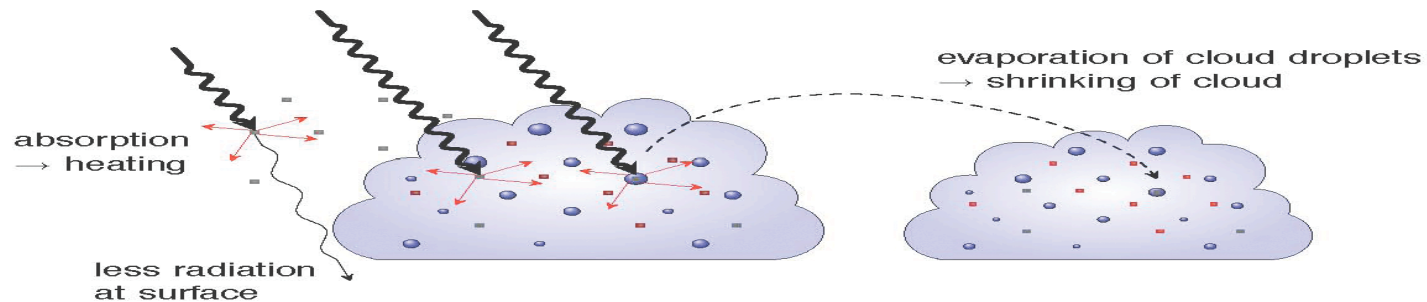
In spring, temperatures begin to rise, the ice evaporates, and the ozone layer starts to recover.

Role of Particles : Effects on Clouds

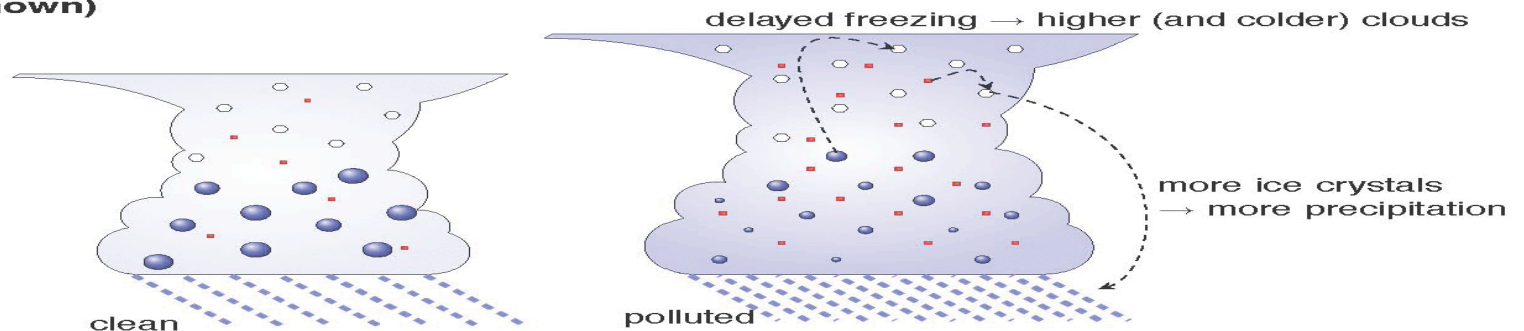
Cloud albedo and lifetime effect (negative radiative effect for warm clouds at TOA; less precipitation and less solar radiation at the surface)



Semi-direct effect (positive radiative effect at TOA for soot inside clouds, negative for soot above clouds)



Glaciation effect (positive radiative effect at TOA and more precipitation), thermodynamic effect (sign of radiative effect and change in precipitation not yet known)



Source: IPCC 2007

Schematic of Aerosol Effects

